# WATER AVAILABILITY AND QUALITY FROM THE STRATIFIED DRIFT IN ANGUILLA BROOK BASIN, STONINGTON AND NORTH STONINGTON, CONNECTICUT

By James W. Bingham

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Prepared in cooperation with the TOWN OF STONINGTON, CONNECTICUT



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# CONVERSION TABLE AND ABBREVIATIONS

For the benefit of readers who prefer metric (International System) units rather than the inch-pound units in this report, the following conversion factors may be used:

Multiply inch- pound unit	by	To obtain metric unit	
	Length		
inch (in.)	25.4	millimeter (mm)	
foot (ft)	0.3048	meter (m)	
mile (mi)	1.609	kilometer (km)	
	Area		
square mile (mi²)	2.590	square kilometer (km²)	
	<u>Flow</u>		
cubic foot per second (ft³/s)	0.02832	cubic meter per second (m³/s)	
cubic foot per second per square mile [(ft³/s)/mi²]	0.01093	<pre>cubic meter per second   per square kilometer [(m³/s)/km²]</pre>	
gallon per minute (gal/min)	0.06308	liter per second (L/s)	
gallon per day (gal/d)	0.003785	cubic meter per day (m³/d)	
million gallons per day (Mgal/d)	3,785	cubic meters per day (m³/d)	
million gallons per day per square mile [(Mgal/d)/mi²]	1,460	<pre>cubic meters per day  per square kilometer [(m³/d)/km²]</pre>	

# Hydraulic Conductivity

hydraulic conductivity,
foot per day (ft/d)

0.3048

hydraulic conductivity,
meter per day (m/d)

**Transmissivity** 

foot squared per day  $(ft^2/d)$ 

0.09290

meter squared per day

 $(m^2/d)$ 

Temperature

degree Fahrenheit (°F)

 $^{\circ}C = 5/9 \times (^{\circ}F - 32)$ 

degree Celsius (°C)

Sea <u>level</u>: In this report "sea level" refers to the National Geodetic

Vertical Datum of 1929 (NGVD of 1929) -- geodetic datum derived from a
general adjustment of the first-order level nets of both the United

States and Canada, formerly called "Sea Level Datum of 1929".

# WATER AVAILABILITY AND QUALITY FROM THE STRATIFIED DRIFT IN ANGUILLA BROOK BASIN, STONINGTON AND NORTH STONINGTON, CONNECTICUT

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#### **ABSTRACT**

The valley of Anguilla Brook is underlain by saturated stratified-drift deposits that, where thick and transmissive, have the potential to yield large quantities of ground water. These deposits are collectively termed the Anguilla Brook aquifer. Long-term yields of four subareas within this aquifer are estimated to range from less than 0.3 to 1.0 million gallons per day. The total yield of all four subareas is estimated to be 2.6 million gallons per day. These yield estimates are based on using the 90-percent duration flow of Anguilla Brook as an index of the water potentially available and on maximum sustainable pumping rates calculated by a mathematical model that used the Theis nonequilibrium equation and image well theory. Development of one or more subareas assumes that most ground water would be derived from induced recharge. This would reduce the flow of Anguilla Brook, and the effect will be most significant during periods when streamflow is low.

Limited sampling and analysis indicate that the quality of both surface and ground water in the Anguilla Brook basin is excellent. The concentrations of all constituents analyzed, with the exception of dissolved manganese and iron, were below the drinking-water limits established by the State of Connecticut, or recommended by the U.S. Environmental Protection Agency.

#### INTRODUCTION

# Purpose and Scope

Saturated deposits of stratified drift drained by Anguilla Brook and its tributaries comprise the Anguilla Brook aquifer, which has the potential for large-scale ground-water development in the Stonington area. The Anguilla Brook aquifer is the only major ground-water source, other than the stratified drift adjacent to the Pawcatuck River, that is within reasonable distance of Pawcatuck, a community of 7,400 that is presently served by public water supply from Westerly, Rhode Island. Planners and town officials are considering the Anguilla Brook aquifer as a possible source of long-term water supply to meet present and future needs of the region.

This report provides estimates of the amounts of ground water available from the Anguilla Brook aquifer and assesses the present quality of surface water and ground water in the area. Four subareas of the Anguilla Brook aquifer have been evaluated to determine their long-term ground-water yield. The analytical technique used in these evaluations first estimates the total amount of ground water potentially available in an area and then employs a mathematical model to determine how much of that total amount can be withdrawn over the long term without excessive drawdowns in the pumping wells. Water samples from Anguilla Brook and the Anguilla Brook aquifer were collected and analyzed to evaluate the present quality of surface water and ground water with respect to drinking-water standards established by the State of Connecticut (Connecticut General Assembly, 1975).

# Physical Setting

The Anguilla Brook basin is located in southeastern Connecticut, in the towns of Stonington and North Stonington (fig. 1). It has an area of about 10 mi² (square miles) and lies midway between the communities of Pawcatuck and Stonington. Land use is mostly rural and the basin is traversed by Interstate Route 95, U.S. Route 1, Connecticut Route 184, and an Amtrak rail line. In 1982, the populations of Stonington and North Stonington were 16,580 and 4,270, respectively (Connecticut Secretary of the State, 1982). An estimated 2,000 persons live in the Anguilla Brook basin.

The climate is moderated by the proximity of Fishers Island Sound but the area is frequently visited by winter and summer continental air masses (Brumbach, 1965). The mean annual air temperature is about 51 °F (degrees Fahrenheit). Mean monthly air temperatures range from a low in January of 22 °F to a high in July of 81 °F (Smith, 1974). The frost-free season extends, on the average, from April 15 to October 25 and results in an average growing season of 193 days. Mean annual precipitation at Groton, Connecticut (the nearest long-term weather station) is 47.51 inches for the 1931-81 period of record.

# Previous Investigations

The ground-water resources of the Anguilla Brook area were briefly discussed by Thomas and others (1968), who used ground-water outflow estimates to conclude that the stratified-drift aquifer might yield as much as 3.4 Mgal/d (million gallons per day) over the long term. Surficial geology of the area was mapped and briefly discussed by Flint (1930). Detailed surficial maps of the eastern and southern parts of the Anguilla Brook basin that lie in the Ashaway and Mystic 7.5-minute quadrangles have been published by the U.S. Geological Survey (Schafer, 1968; Upson, 1971). Mapping in the western part of the area (Old Mystic 7.5-minute quadrangle) is in progress.

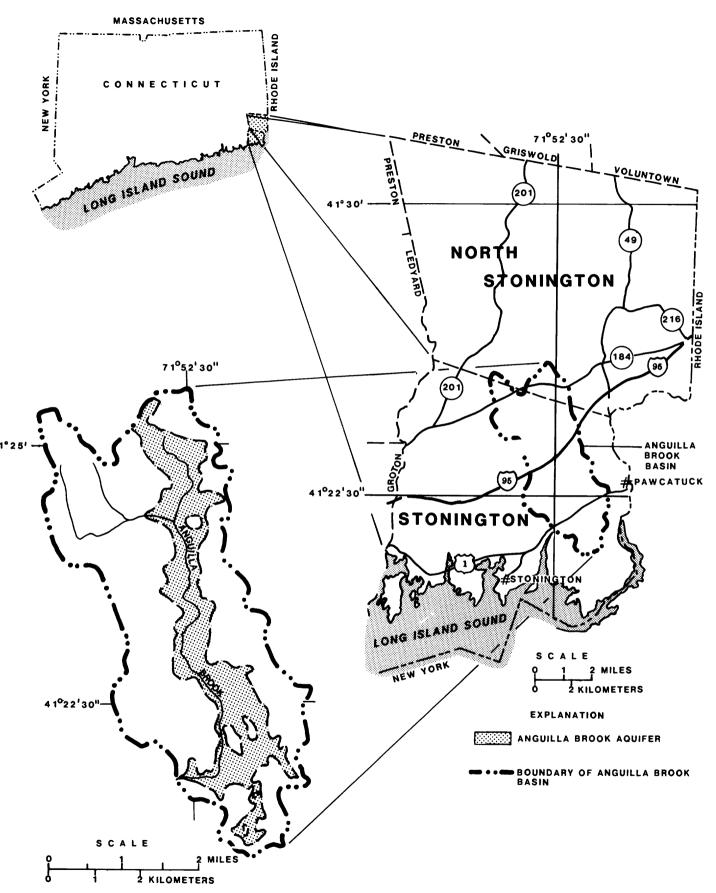


Figure 1.--Location of Anguilla Brook basin and Anguilla Brook aquifer, Stonington and North Stonington, Connecticut.

# Acknowledgments

This investigation was conducted by the U.S. Geological Survey in cooperation with the Town of Stonington, Connecticut. The author is grateful for the assistance provided by officials of Stonington and concerned citizens and property owners in Stonington and North Stonington; their cooperation and contributions are sincerely appreciated.

# DATA COLLECTION

During this investigation, a variety of hydrogeologic data were collected in the Anguilla Brook basin. The data are the basis for evaluating the long-term yield of the Anguilla Brook aquifer and assessing present water-quality conditions.

Seismic-refraction surveys were run along six profile lines that extend across the Anguilla Brook aquifer in order to determine depths to bedrock and depths to the water table. The seismic information was also used to design a subsequent test-drilling program. Interpretive hydrogeologic cross-sections along the six survey lines are shown in figure 15 (located at the end of this report). Fifteen test borings were made with a hollow-stem power auger to determine or verify depths to bedrock and depths to the water table and to collect samples of stratified drift for grain-size analysis. At nine of these sites, observation wells with 2-inch diameter PVC (polyvinyl chloride) plastic screens were installed to collect ground-water samples for chemical analysis and to periodically measure ground-water levels. Descriptions of these wells and test holes are given in table 8 (located at the end of this report); descriptions of other wells and test holes that provided additional hydrogeologic information are also given in this table.

The earth materials encountered while drilling the wells and test holes are described in table 9 (located at the end of this report). Selected samples of stratified drift, collected during the test drilling phase, were dried and sieved in order to determine the grain-size distribution of the material. As noted in a later section of this report, this information is used to estimate some of the hydrogeologic characteristics of stratified drift. The results of the grain-size analyses are given in table 10 (located at the end of this report).

Locations of the seismic-refraction profile lines, wells, test holes, streamflow measurement sites, and water-quality sampling sites are shown on plate 1.

#### DATA ANALYSIS

The first phase of data analysis consisted of integrating information on depths to the water table and bedrock with average hydraulic conductivity (estimated from grain-size characteristics and lithologic descriptions of stratified drift) to produce maps of saturated thickness and transmissivity of the Anguilla Brook aquifer. Subareas were selected from these maps that had relatively high saturated thickness and average transmissivity. The four subareas shown in figure 10 were selected for evaluation of their long-term yields.

The subarea evaluation consisted of estimating the water potentially available. Each subarea is traversed by Anguilla Brook and the water potentially available from this source is much greater than the recharge from precipitation. For this reason, and to simplify the evaluation, a low-flow characteristic (the 90-percent duration flow) of Anguilla Brook was selected as the water potentially available over the long term. The criteria for selection and computation process are discussed in more detail in the section of this report entitled "Ground-Water Potentially Available". Maximum pumping rates that could be sustained in each subarea on the basis of the hydrogeologic characteristics and assumed boundary conditions of the aquifer were computed by a simple mathematical model. The model elements and assumptions and the modeling procedure are outlined in the section of this report entitled "Maximum Long-Term Pumping Rates".

Long-term ground-water yields for each subarea are estimated for two different conditions. The first condition assumes no development of any upstream areas and the water potentially available is compared to the maximum pumping rates and the lesser quantity is considered to be the long-term ground-water yield. The second condition assumes maximum development of all upstream subareas and the water potentially available is adjusted to reflect the upstream withdrawals. The maximum pumping rates are then compared to these adjusted values and the lesser quantity is the long-term ground-water yield.

#### BASIN HYDROGEOLOGY

# Geologic Framework

Stratified drift comprises the principal aquifer in the Anguilla Brook basin and is the focus of this study. This material was deposited by meltwater streams flowing southward away from the margin of glacial ice that stood against the low ridge located between Rocky Hollow Road and Assekonk Swamp in North Stonington (plate 1). The materials that make up the stratified-drift deposits were sorted by the actions of the meltwater streams—coarse—grained material was deposited in the northern part of the valley and finer—grained material was deposited to the south. In the southern part of the valley, the fine—grained sediments are overlain by several feet of coarse material. The geologic conditions that produced this "gravel cap" are unknown. As the ice mass continued its northward retreat, a drainageway opened up down the Shunock River valley and diverted meltwaters away from the Anguilla Brook valley.

Till and bedrock are the other major geologic units in the Anguilla Brook basin. These units underlie and are adjacent to the stratified-drift deposits as shown in figure 6, and are considered to form relatively impermeable boundaries for the Anguilla Brook aquifer. Wells tapping till and bedrock are much less productive than those in the stratified drift, but may yield water supplies adequate for homes, farms, or small businesses. Yields typically range from 1 to 10 gal/min (gallons per minute)--an amount too small for public-supply purposes.

Till is composed of unconsolidated and unsorted glacial material deposited near the margin of an ice sheet. In Connecticut, it commonly forms a layer of debris, less than 15 feet thick, that mantles much of the bedrock surface. This till has a low permeability in comparison to stratified drift and can be considered an effective barrier that significantly limits the flow of ground water between these two units.

Bedrock underlies both till and stratified drift in Stonington and North Stonington. It consists of metamorphic rock, generally gneiss (Rodgers, 1982). Bedrock also has a relatively low permeability, except where networks of open joints and fractures provide conduits for ground-water movement. Wells drilled into bedrock generally intercept these joints and fractures. Depending in part on the number and size of fractures encountered, the yield from a drilled well tapping bedrock in the Anguilla Brook area averages about 5 gal/min, and rarely exceeds 15 gal/min.

#### Circulation of Water

All the water in Anguilla Brook basin is derived from precipitation falling on the land surface. This water is constantly in motion. Some moves overland to streams (Wheeler and Anguilla Brooks and their tributaries) and leaves the basin as a component of streamflow; some moves downward through the soil, recharges the stratified-drift aquifer, and then leaves the basin either as ground-water runoff (another component of streamflow), or underflow; and some is stored temporarily on the land surface or in the soil zone and leaves by means of evapotranspiration. Water that is lost from the basin is eventually replenished by precipitation.

The hydrologic system in the Anguilla Brook basin is in balance and water entering, stored within, and leaving can be accounted for. This balance is represented by the equation:

```
P = SW(ro) + GW(ro) + U + ET ± S
where:
    P = Precipitation,
    SW(ro) = Surface-water runoff (streamflow component),
    GW(ro) = Ground-water runoff (streamflow component),
    U = Underflow,
    ET = Evapotranspiration, and
    S = Changes in storage
```

# Precipitation

Precipitation in the study area occurs throughout the year and monthly averages range from 2.5 to 5 inches. Figure 2 shows average monthly precipitation at nearby Groton, Connecticut (National Weather Service Index Number 3207), for the 1941-70 period. Annual precipitation at Groton for 1931-81 is shown in figure 3. During this 51-year period, precipitation averaged 47.51 inches annually. Figure 3 also includes a 5-year moving average that shows dry cycles, indicated by a declining line (1940-47, 1962-66) and wet cycles, indicated by a rising line (1947-55, 1966-73).

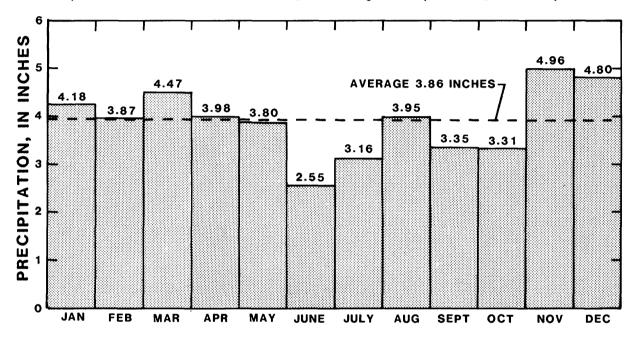


Figure 2.--Monthly average precipitation at Groton, Connecticut, 1941-70.

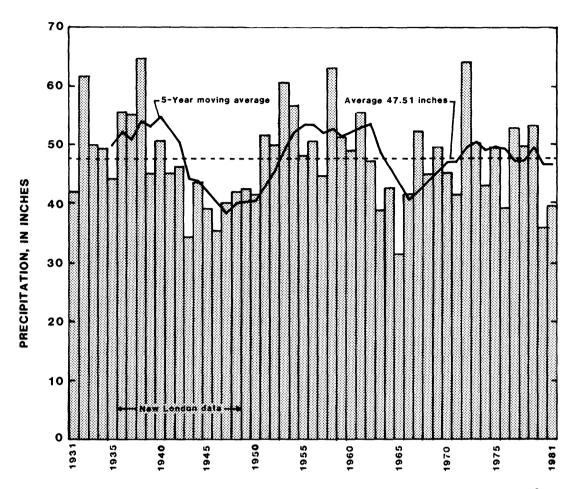
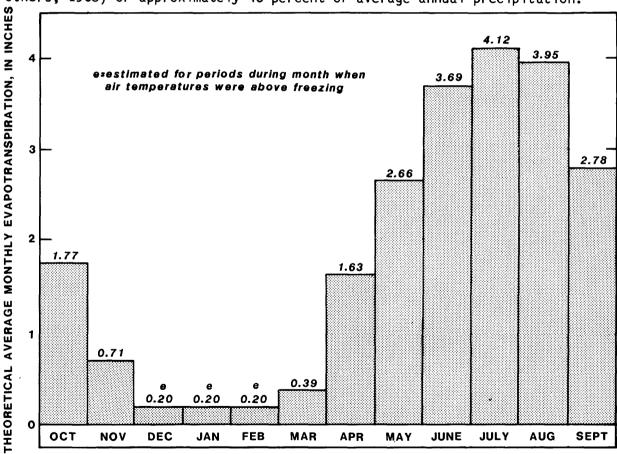


Figure 3.--Annual precipitation and 5-year moving average of annual values at Groton, Connecticut, 1931-81.

# Evapotranspiration

Evapotranspiration is the direct loss of water to the atmosphere by evaporation from water surfaces and moist soil, and transpiration from living plants. Evapotranspiration is seasonal; plants are most active and may transpire large amounts of water during the frost-free growing season. Some plants such as evergreens also transpire during the winter months but at a reduced rate. Evaporation occurs all year long but is greatest during the summer months. These seasonal fluctuations are illustrated in figure 4, which shows how the theoretical monthly average evapotranspiration ranges from 0.20 inches in the winter months (December, January, and February) to 4.12 inches in the summer (July). The values in figure 4 are computed by a method similar to that of Thornwaite and Mather (1957) and are adjusted to account for changes in storage.

The estimated average annual evapotranspiration in the Anguilla Brook area during the 1931-60 period of record was about 22.3 inches (Thomas and others, 1968) or approximately 46 percent of average annual precipitation.



(From Thomas and others, 1968, fig.6)
Figure 4.--Estimated average monthly evapotranspiration in southeastern
Connecticut, 1931-60.

#### Streamflow

Streamflow in the Anguilla Brook basin is estimated to have averaged about 2.0 (ft<sup>3</sup>/s)/mi<sup>2</sup> (cubic feet per second per square mile) or about 12.8 Mgal/d during the 1931-60 period (Cervione and others, 1968). Long-term, average streamflow values include flood flows. For water-supply planning purposes, low-flow values are needed that reflect conditions during dry Indices of low flow that are commonly used for water-resources planning in Connecticut are (1) the average streamflow equaled or exceeded 90 percent of the time or its approximate equivalent, the 30-day, 2-year low flow, and (2) the average streamflow equaled or exceeded 99 percent of the time or its approximate equivalent, the 7-day, 10-year low flow. Detailed discussion of these low-flow indices and how they are derived are found in Weiss and others (1982). Low-flow characteristics can be determined from records of stream-gaging stations or estimated by various regionalization techniques. The flow of Anguilla Brook has not been gaged and the low-flow values cited in this report are based on regional flow-duration curves developed by Thomas (1966). Adjustments representing drier than average and wetter than average conditions followed the procedures outlined in Weiss and others (1982, p. 12-17).

Figure 5 is a flow-duration curve for Anguilla Brook, for the period 1931-60 prepared by regionalization methods. It shows flows that are equaled or exceeded for various percentages of time in average, wettest, and driest years. The 90-percent duration flow of Anguilla Brook above Green Haven Road ranges from 1.5 Mgal/d during the driest year to about 5.0 Mgal/d during the wettest year. The average 90-percent duration flow at this site is about 2.5 Mgal/d.

Table 1, a summary of the low-flow characteristics of Anguilla Brook, lists the annual lowest mean flows for 2- and 10-year recurrence intervals that will occur for time periods ranging from 3 to 365 consecutive days. As noted above, the 90-percent duration flow value is approximately equal to the 30-day, 2-year low flow.

Low flows of Anguilla Brook were measured directly at four stations on August 19, 1982. The locations of these streamflow-measurement sites are shown on plate 1. The flow at nearby Pendleton Hill Brook, a long-term gaging station (station 01118300) was 0.86 ft $^3$ /s (cubic feet per second) on this date; this was approximately equal to the 84-percent duration flow. The measured flow of Anguilla Brook ranged from 0.11 ft $^3$ /s at Minor Pentway (station 01118530) to 2.5 ft $^3$ /s at Connecticut Route 1 (station 01118550). The flows at Nathan Wheeler farm (station 01118535) and South Anguilla Road (station 01118548) were 1.4 ft $^3$ /s and 2.1 ft $^3$ /s respectively.

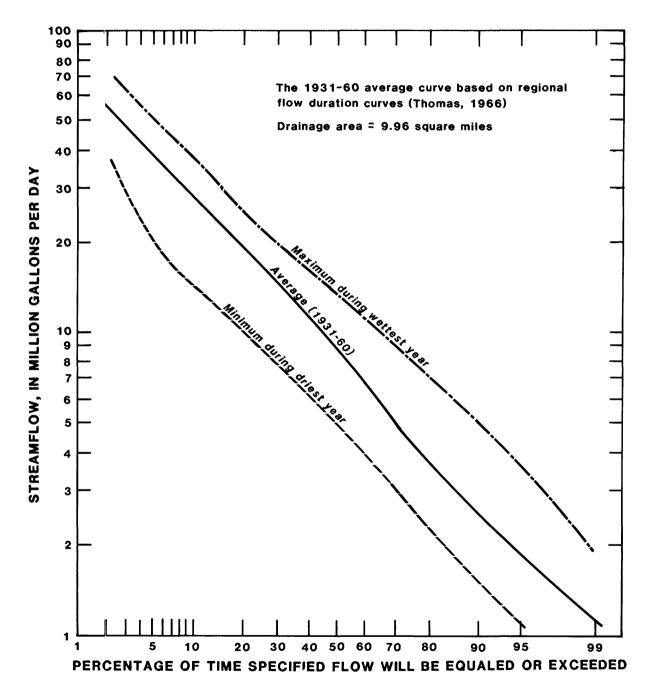


Figure 5.--Duration of daily flows of Anguilla Brook above Green Haven Road, 1931-60.

Table 1.--Annual lowest mean flows for Anguilla Brook above Green Haven Road [All flows are in million gallons per day]

Recurrence			Conse	cutive da	ays		
interval (years)	3	7	30	60	120	183	365
2	1.6	2.0	2.6	3.2	4.1	5.4	11.6
10	1.1	1.4	1.8	1.9	2.6	3.2	5.8

# Ground-Water Recharge

The precipitation that does not flow directly over the land surface to streams and is not evapotranspired, percolates downward and recharges the aquifers in Anguilla Brook basin. Measurements of ground-water recharge are difficult to obtain and none are available in Connecticut. However, recharge can be estimated over a period of time by subtracting surface runoff and total evapotranspiration from precipitation. This technique assumes that the amount of ground water in storage at the beginning and end of the specified time is the same. Otherwise, the differences in ground-water storage must be taken into account. In southeastern Connecticut during the 1931-60 period, annual precipitation averaged 48.2 inches, surface runoff averaged 14.8 inches, and evapotranspiration averaged 22.3 inches (Thomas and others, 1968). These figures suggest that an average recharge of about 11.1 inches per year from precipitation is reasonable for the Anguilla Brook basin. Recharge to the stratified-drift aquifer is somewhat higher and is estimated to average about 20 inches per year.

Under conditions of development, induced recharge from streams can significantly augment recharge from precipitation. Pumping wells near a stream often cause the water table to fall below the level of the stream. This ultimately results in the infiltration of water from the stream into the adjacent aquifer. The amount of water available for this induced recharge, on a sustained basis, is limited by the flow of the stream.

#### HYDROGEOLOGIC CHARACTERISTICS OF STRATIFIED DRIFT

The amount of ground water that can be developed from a stratified-drift aquifer is affected by many factors. Of major importance are the hydrogeologic characteristics: saturated thickness, hydraulic conductivity, transmissivity (the product of average hydraulic conductivity and saturated thickness), and specific yield. Test holes, stratified-drift samples from these holes, water-level measurements, and seismic-refraction profiles provided information about these characteristics that are discussed in this section.

The saturated thickness of stratified drift that forms the Anguilla Brook aquifer is equal to the vertical distance from the water table to the underlying till or bedrock surface (fig. 6). It was determined by seismic-refraction profiling, test drilling, and evaluating information from existing wells. The locations of these data sites are shown on plate 1. The saturated thickness of the four subareas of the aquifer where long-term ground-water yields have been estimated is shown in figures 11 to 14.

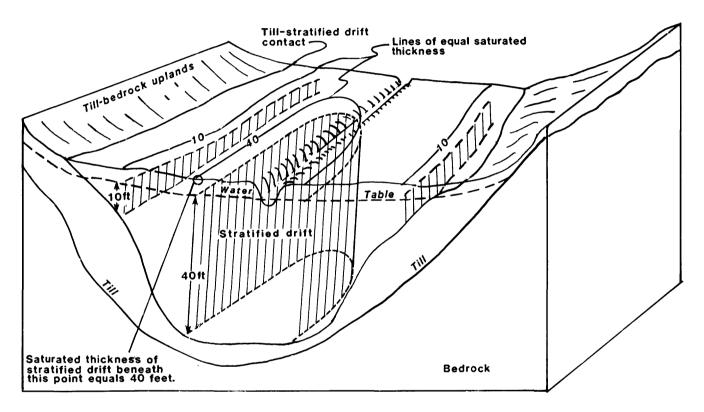
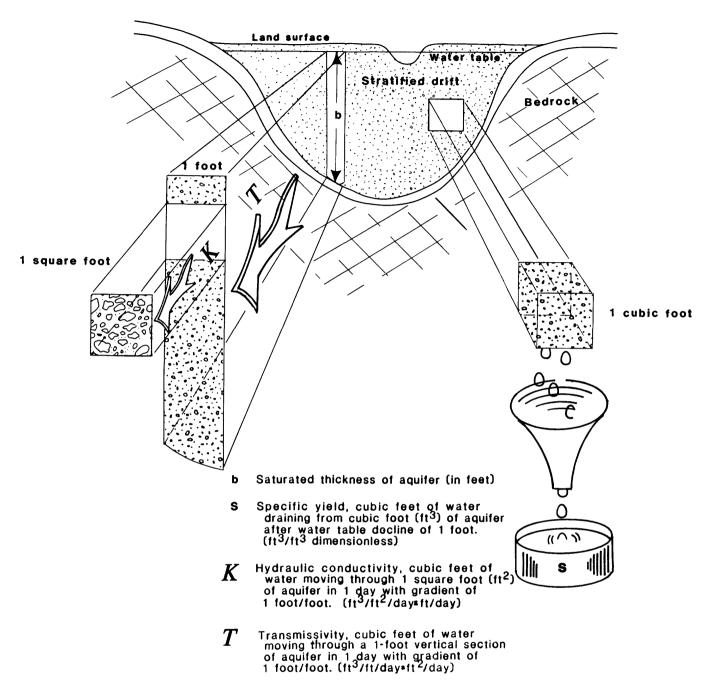


Figure 6.--Spatial relations between a stratified-drift aquifer and adjacent till and bedrock and the saturated thickness of the stratified drift.

The hydraulic conductivity of stratified drift is a measure of its ability to transmit water. In English units, a material would have a hydraulic conductivity of 1 foot per day if it transmitted 1 cubic foot of water in 1 day's time through a cross-sectional area of 1 square foot. The water is assumed to be at field viscosity, and to be under a hydraulic gradient of 1-foot change in head over 1-foot in length of flow path. Hydraulic conductivity is illustrated in figure 7. In the Anguilla Brook aquifer, the estimated hydraulic conductivity of the coarse-grained stratified drift generally ranges from 50 to 200 ft/d (feet per day).



(from Weiss and others, 1982)

Figure 7.--Hydraulic characteristics of a stratified-drift aquifer.

In this investigation, hydraulic conductivities are based on relations between this characteristic and the median-grain size and sorting of the stratified drift developed in previous Connecticut studies and summarized by Ryder and others (1970). The theoretical basis for these relations is discussed by Krumbein and Monk (1942), and Masch and Denny (1966). Estimates of the average hydraulic conductivity at different sites were developed from logs of test holes that have grain-size analyses of samples of the materials penetrated.

The transmissivity of an unconfined stratified-drift aquifer is a measure of the rate at which a fluid will pass through a section of it. In English units, an aquifer would have a transmissivity of 1 square foot per day if it transmitted 1 cubic foot of water through a section 1 foot in width that extended from the water table to the bottom of the aquifer. The water is assumed to be at field viscosity and to be under a hydraulic gradient of 1-foot change in head over 1-foot length of flow path. Transmissivity is equivalent to the average hydraulic conductivity times the saturated thickness and is illustrated in figure 7. In the Anguilla Brook aquifer, estimated transmissivity generally ranges from 2,000 to 6,000 ft $^2$ /d (feet squared per day).

Transmissivities for the Anguilla Brook aquifer are based on the average hydraulic conductivity values determined for samples of stratified drift from test holes and information about changes in saturated thickness. Transmissivity was estimated at test-hole sites by summing the individual hydraulic conductivities of each foot of aquifer over the total saturated thickness. The technique is described by Weiss and others (1982). Zones of equal transmissivity for the four subareas of the Anguilla Brook aquifer evaluated for long-term yield were mapped by interpolating between data points. The resulting transmissivity distributions are shown in figures 11 through 14.

Specific yield is the ratio of the volume of water that a saturated earth material will yield, by gravity drainage, to its own volume. It is a term in flow equations used for calculating water-table drawdowns over different time periods in unconfined aquifers and is analogous to the storage coefficient of confined aquifers. Specific yield is controlled, in part, by the number and size of the intergranular pore spaces of the aquifer materials. Figure 7 illustrates the concept of specific yield and shows it is a dimensionless value obtained by measuring the volume of water drained by gravity, from a known volume of aquifer, over a relatively long period of time (100 to 200 days).

In the Anguilla Brook aquifer, specific yield was not measured directly. It is estimated to range from 0.15 to 0.30 and average 0.20, based on studies of similar unconsolidated sediments conducted elsewhere.

#### GROUND-WATER AVAILABILITY

# Ground Water Potentially Available

The ground water potentially available for development from the Anguilla Brook aquifer, over the long term, is assumed to be equal to the 90-percent duration flow of Anguilla Brook (or this value adjusted for the effects of upstream development) at the point of anticipated withdrawal. This particular flow is commonly used in evaluating the yields of stratified-drift aquifers in Connecticut (Cervione and others, 1972; Mazzaferro and others, 1979) as it represents a fairly dependable quantity available for induced recharge 90 percent of the time. Selection of a higher flow could result in higher estimated long-term yields but would also result in potentially greater reductions in streamflow, particularly in summer months when streamflow is low and ground-water withdrawals are high. In fact, if all of the 90-percent flow was to infiltrate into the aquifer, all or parts of Anguilla Brook would be dry 10 percent of the time.

While the 90-percent duration streamflow is used as a limiting factor for estimating long-term yields, it is unlikely that all the water would come from the stream; some would be derived from recharge from precipitation and depletion of storage in the aquifer. This means that (1) if all or parts of Anguilla Brook are allowed to go dry 10 percent of the time, long-term yields based on the 90-percent duration flows are conservative, and (2) if withdrawals are limited to this flow value, all or parts of Anguilla Brook would go dry less than 10 percent of the time.

The water assumed to be potentially available for development in the four subareas over the long term is computed using a unit area 90-percent duration flow of 0.4 ( $ft^3/s$ )/ $mi^2$  [0.26 (Mgal/d)/ $mi^2$  (million gallons per day per square mile)]. This is the estimated average value for the entire basin as determined from figure 5. Estimated amounts of water potentially available over the long term for the four subareas and their upstream drainage areas are given in tables 3 and 4.

# Maximum Long-Term Pumping Rates

In southern New England, long-term pumping rates and yields of stratified-drift aquifers have been estimated by mathematical models that use the Theis nonequilibrium equation (Theis, 1935) to describe flow, and the theory of image wells (Ferris and others, 1962) to simulate hydraulic boundary conditions (Allen and others, 1966; Rosenshein and others, 1968; Ryder and others, 1970). The model analysis requires hydrologic information about the aquifer under investigation (saturated thickness, average transmissivity, average specific yield, location and type of hydraulic boundaries), and physical information about the method and pattern of withdrawals (number and spacings of hypothetical pumping wells, well-construction characteristics, period of pumping).

The modeling procedure for calculating maximum long-term pumping rates in the Anguilla Brook aquifer is the same as in previous Connecticut studies and is described by Cervione and others (1972) and Mazzaferro and others (1979). To facilitate the analysis, a Fortran-language program was used that allows a computer to calculate the drawdowns and yields of as many as 16 pumping wells. The calculation process also adjusts the drawdowns and yields to account for the effects of dewatering of the aquifer, partial penetration of each well, pumping of nearby wells, and hydraulic boundaries.

After selecting initial pumping rates for the wells, the adjusted drawdowns are calculated at each well. The maximum pumping rate is specified at the rate that results in an adjusted drawdown that is about 1 foot above the top of the well screen. Pumping rates are incrementally increased or decreased at each well until the adjusted drawdown meets the criteria (1 foot above the top of the well screen) for all wells in the modeled area. The following assumptions regarding aquifer and well characteristics and period of pumping are incorporated into the models of each of the four subareas.

- (1) Specific yield of the aguifer is equal to 0.20;
- (2) Maximum available drawdown, at each well, is 70 percent of the saturated thickness;
- (3) Screen length for each well is 30 percent of the saturated thickness;
- (4) Effective well radius is 2 feet;
- (5) Pumping period is 180 days, and there is no recharge during this period (this is approximately equal to the length of time that there is little or no recharge from precipitation each year and if pumpage can be sustained over this period, it is likely the same rate can be sustained indefinitely);
- (6) Wells are 100-percent efficient;
- (7) Vertical-to-horizontal hydraulic conductivity ratio (K<sub>V</sub>:K<sub>h</sub>) is 0.10;
- (8) Wells are generally located in the thickest, most transmissive parts of the aguifer.

Models of the four subareas of the Anguilla Brook aquifer analyzed during the course of this investigation have two to six hypothetical pumping wells, generally located in the most favorable sites. The number of hypothetical wells is partly controlled by the transmissivity and saturated-thickness distribution in each subarea but they are also spaced at distances ranging from 750 to 1,800 feet to reduce the effects of well interference. Locations of the hypothetical pumping wells are shown in figures 11 to 14.

One of the assumptions of the nonequilibrium equation (Theis, 1935), the basis of this method of analysis, is that the aquifer is of infinite areal extent. The Anguilla Brook aquifer is finite--limited by natural features that also form hydraulic boundaries. Two different boundary conditions are considered in the mathematical models of the four subareas. They are impermeable-barrier or no-flow boundaries and line-source or constant head boundaries.

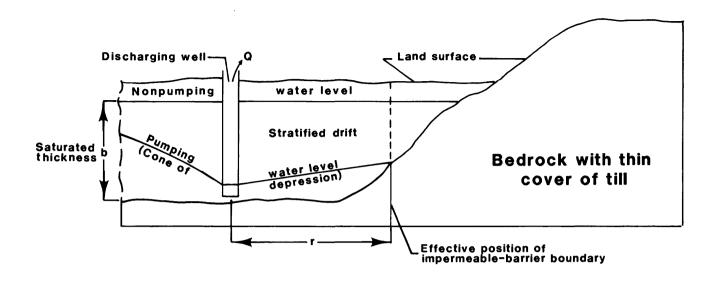
The first boundary type is formed by relatively impermeable materials or features that are in contact with the aquifer, and little or no ground water flows across the contact (eg., adjacent till, bedrock, and thin, fine-grained stratified drift). The second boundary type is formed by streams or other surface-water bodies that contribute large amounts of water to the aquifer. Both types of hydraulic boundaries are illustrated in figure 8.

In the models of the four subareas, impermeable-barrier boundaries are idealized as straight, vertical planes placed to approximately coincide with the 10-foot saturated thickness contour lines along the aquifer margins. Some ground water may flow into the aquifer across these planes which would tend to make the maximum long-term pumping rates derived by the models more conservative. Line-source boundaries are also idealized as vertical planes and, in most place, are located along the axis of Anguilla Brook. They represent a constant source of water recharging the aquifer. In theory, the water level is constant along the line-source boundary and the effects of pumping are not felt across the vertical plane. Field conditions differ from ideal in that Anguilla Brook is relatively small and does not fully penetrate the aquifer (fig. 8). Consequently, large pumpage on one side of the stream would likely affect water levels in the stream and in the aquifer on the other side.

Image wells are used in the mathematical model to simulate the effects of the specified boundary conditions as shown in figure 9. Image wells may be either discharging or recharging, depending on the boundary condition they are simulating. When more than one boundary is involved, it is necessary to account for the multiple boundaries using repeated image well arrays (Ferris and others, 1962, p. 156). In theory, these image wells can be repeated infinitely, but in this study they are terminated after 13 repetitions or when the drawdown or buildup of the water table due to the image well is less than 0.00001 foot.

Well and aguifer characteristics, assumed boundary conditions and estimated maximum long-term pumping rates for the four subareas shown in figure 10 are summarized in table 2. In three of the four subareas, two different boundary conditions are incorporated into and evaluated by the mathematical model. The first assumed that two or three sides of the subarea act as impermeable-barrier boundaries. The second assumed that one or two sides of the subarea act as impermeable-barrier boundaries and one side acts as a line-source boundary. It is also assumed that the remaining one or two sides have neither type of boundary and the aquifer is continuous in those directions. These alternative boundary configurations result in two values of estimated, maximum long-term pumpage for three of the subareas. The smaller value represents the long-term pumpage if only impermeable-barrier boundaries are effective. The larger value represents a more optimistic situation where Anguilla Brook acts as a line-source boundary. In the northern subarea, only one boundary configuration (three impermeable-barrier boundaries and one side open or continuous) was evaluated because the size and flow of Anguilla Brook at this point are so small it is unlikely to constitute a line-source boundary. Consequently, only one estimate of maximum long-term pumping rate for this subarea is listed in tables 2, 3, and 4.

# A. IMPERMEABLE-BARRIER BOUNDARY



# **B. LINE-SOURCE BOUNDARY**

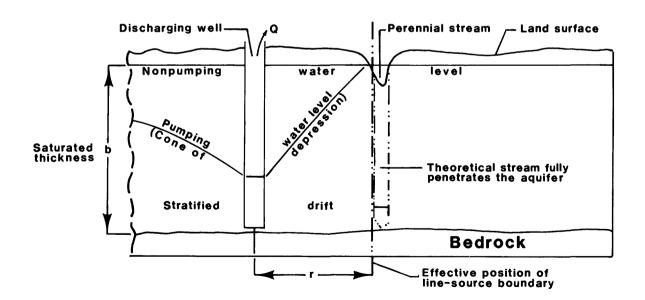
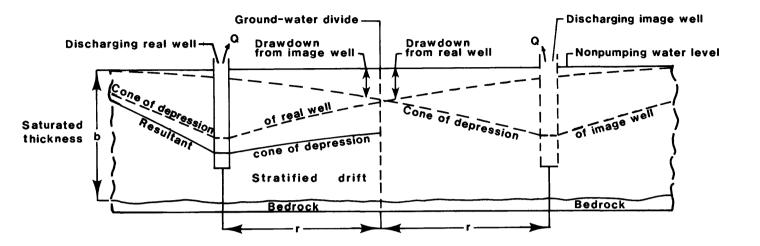


Figure 8.--Cross sections of a stratified-drift aquifer showing hydraulic boundaries and their effect on pumping water levels. (Modified from Ferris and others, 1962, figs. 35, 37.)

A. Discharging image well simulating the effects of an impermeable-barrier boundary.

The hydraulic conditions in the infinite aquifer shown below are the same as in figure 8A.

The ground-water divide across which there is no flow is in the same location as the impermeable-barrier boundary in figure 8A.



B. Recharging image well simulating the effects of a line-source boundary.

The hydraulic conditions in the infinite aquifer shown below are the same as in figure 8B.

The line of zero drawdown (constant head) is in the same location as the line-source boundary in figure 8B.

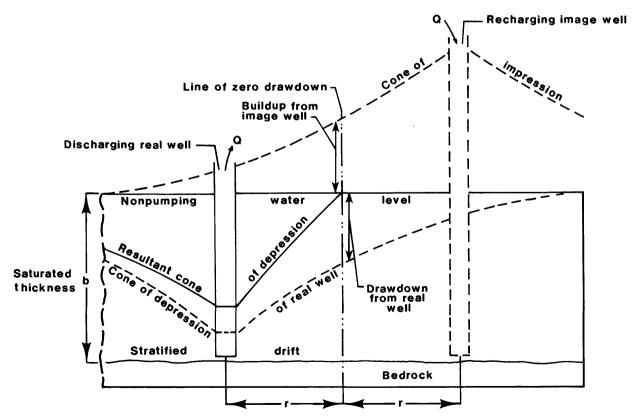


Figure 9.--Cross sections showing how image wells are used to simulate the hydraulic effects of boundaries. (Modified from Ferris and others, 1962, figs. 35, 37.)

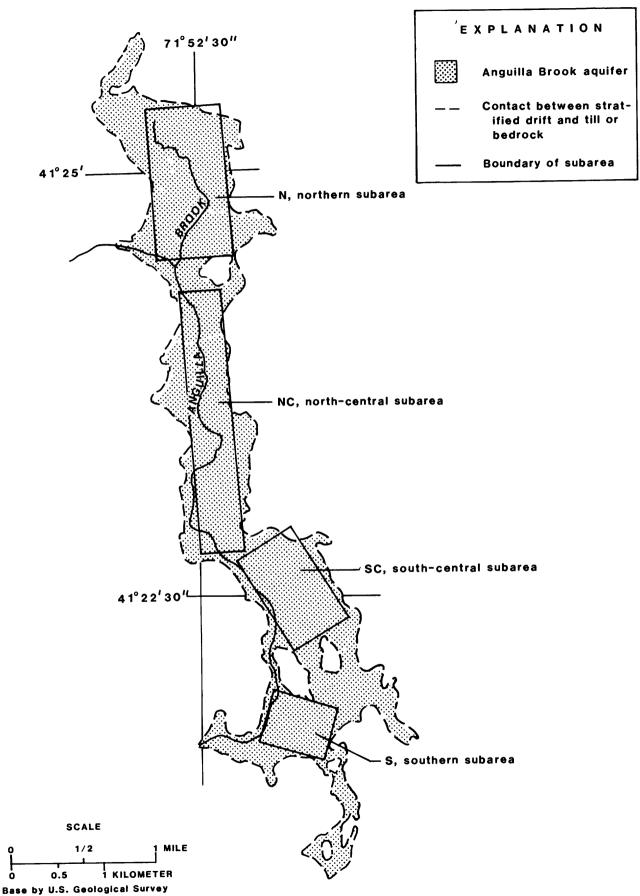


Figure 10.--Locations of favorable subareas of the Anguilla Brook aquifer.

Table 2.--Boundary conditions, selected well and aquifer characteristics, and estimated maximum long-term pumping rates for four subareas of the Anguilla Brook aquifer [Position of boundary indicated by letter: W, west; E, east; N, north; S, south.

Type of boundary indicated by letter: B, impermeable-barrier boundary; L, line-source boundary; O, no boundary. K, ratio of vertical to horizontal hydraulic conductivity]

					Average		Maximum lo pumping r	-
Name and map symbol (See fig. 10	Boundary conditions	Number	Effective well radius	Saturated thickness	aquifer transmissivity (feet squared	K <sub>v</sub> :K <sub>h</sub>	Per well (gallons	All wells (million gallons
for location)	WENS	of wells	(feet)	(feet)	per day)	ratio	per minute)	per day)
Northern	B B B O	5	2.0	40	4,000	0.1	163	1.5
N				40			145	
				50			216	
				40			153	
				75			380	
North-central	L B B O	6	2.0	20	2,000	.1	71	.7
NC				20			72	
				25			80	
				25			79	
				25			81	
				35			115	
North-central	ввво	6	2.0	20	2,000	.1	61	.6
NC				20			54	
				25			65	
				25			57	
				25			56	
				35			98	
South-central	LBOB	4	2.0	30	3,000	.1	148	1.0
SC				35			171	
				40			194	
				30			157	
South-central	ввов	4	2.0	30	3,000	.1	91	.7
SC				35			112	
				40			142	
				30			111	
Southern	L B O O	2	2.0	25	2,500	.1	104	.3
S				25			104	
Southern	B B O O	2	2.0	25	2,500	.1	93	.3
S				25			93	

# Long-Term Ground-Water Yields

The following discussion of long-term ground-water yields from the Anguilla Brook aquifer is based on (1) estimates of the ground water potentially available over the long term, and (2) estimates of the maximum long-term pumping rates. The aquifer is divided into four subareas, each traversed by Anguilla Brook (fig. 10). The ground water potentially available over the long term in each subarea is initially assumed to be equal to the 90-percent duration flow of Anguilla Brook. Water available from recharge from precipitation or storage in the aquifer is assumed to be negligible compared to that from induced recharge from the brook. The maximum long-term pumping rates in each subarea are estimated by use of the previously described mathematical model that is based on the Theis nonequilibrium equation (Theis, 1935) and the theory of image wells (Ferris and others, 1962).

Maximum long-term pumping rates (table 2) can be sustained in each subarea as long as they do not exceed the water potentially available. If this is the case, the long-term ground-water yield is considered to be equal to the maximum pumping rate. Conversely, where the maximum long-term pumping rate is greater than the water potentially available, the long-term ground-water yield is considered to be equal to the water potentially available.

In estimating long-term ground-water yields in all but the northern subarea, two different conditions of development have been considered. first condition assumes there is no development of any of the upstream subareas that would affect the amount of water potentially available. The analysis under this condition compares the ground water potentially available with the maximum pumping rate and considers the smaller of these to be the estimated long term ground-water yield. The results of this evaluation are summarized in table 3. The second condition assumes maximum development of all upstream subareas and that the water withdrawn from these areas is exported from the basin. Consequently, the water potentially available to a downstream subarea (the 90-percent duration flow) is reduced by an equivalent amount. The impacts of upstream development are cumulative and reductions in the amount of water potentially available are greater in each succeeding downstream subarea. The analysis under this condition of development compares the adjusted value of the ground water potentially available and the maximum pumping rate and considers the smaller of these to be the estimated long-term ground-water yield. The results of the evaluation under this development alternative are summarized in table 4.

Table 3.--Long-term ground-water yields in subareas of the Anguilla Brook aquifer with no upstream development

A	В	С	D	E
Subarea name and map symbol	Upstream drainage	Estimated amount of water poten- tially available (equal to 90-percent duration flow of Anguilla Brook)	Maximum long-term pumping rate determined from mathematical models	Long-term ground-water yields (the lesser of column C or D)
(See fig. 10	area	(million gallons	(million gallons	(million gallons
for location)	(square miles)	per day)	per day)	per day)
Northern (N)	3.4	0.9	1.5	0.9
North-central (NC)	7.0	1.8	$\frac{1}{0.6}$ - 0.7	0.6 - 0.7
South-central (SC)	9.6	2.5	$\frac{1}{0.7}$ - 1.0	0.7 - 1.0
Southern (S)	9.9	2.6	0.3	0.3

<sup>&</sup>lt;sup>1/</sup> Smaller number is the rate calculated with only impermeable-barrier boundaries and larger number is the rate calculated with a line-source boundary. See table 2.

#### Northern Subarea

This subarea, located in the upper part of the Anguilla Brook aguifer, contains the thickest and most coarse-grained stratified drift in the basin. Transmissivities generally range from 3,000 to 6,000 ft<sup>2</sup>/d and saturated thickness, in places, exceeds 75 feet. The geologic contact between till and stratified drift, altitude of the bedrock surface, saturated thickness, transmissivity, assumed boundary conditions, and locations of hypothetical pumping wells are shown in figure 11. The aquifer characteristics and other features of the mathematical model of this area are given in table 2. Five hypothetical wells, placed at approximately 1,000-foot intervals, are used to withdraw water for a 180-day period. Only one boundary configuration, consisting of three impermeable-barrier boundaries, was assumed in evaluating this subarea. The maximum, long-term pumping rate calculated for these conditions was 1.5 Mgal/d, nearly 70 percent greater than the 0.9 Mgal/d estimated to be potentially available at this site over the long term. Thus, the long-term ground-water yield in the northern subarea, as indicated in tables 3 and 4, is estimated to be 0.9 Mgal/d, the lesser of these two values.

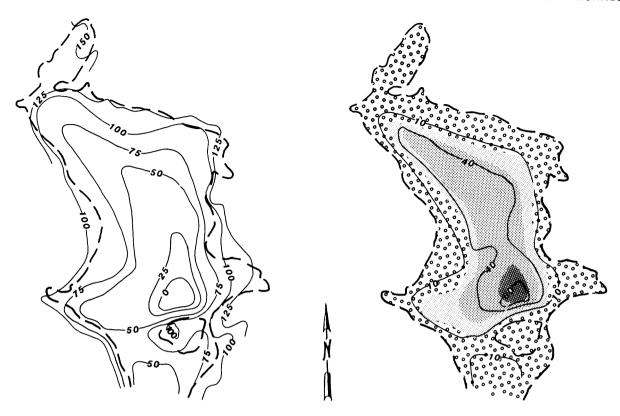
Table 4.--Long-term ground-water yields in subareas of the Anguilla Brook aquifer with maximum upstream development

Α	В	С	D	E	F
Subarea name and map symbol (See fig. 10 for location)	Estimated amount of water poten- tially available (equal to 90-percent duration flow of Anguilla Brook) (million gallons (per day)	Cumulative ground-water withdrawals from upstream subareas (million gallons per day)	Water potentially available adjusted for effects of upstream development (90-percent duration flow of Anguilla Brook minus cumulative withdrawals) (million gallons per day)	Maximum long-term pumping rate determined from mathematical models (million gallons per day)	Long-term ground-water yields (the lesser of column D or E) (million gallons (per day)
Northern (N)	0.9	0	0.9	1.5	0.9
North-central (NC)	1.8	0.9	0.9	$\frac{1}{0.6}$ - 0.7	0.6 - 0.7
South-central (SC)	2.5	1.5 - 1.6	0.9 - 1.0	$\frac{1}{2}$ 0.7 - 1.0	0.7 - 1.0
Southern (S)	2.6	2.2 - 2.6	0 - 0.4	0.3	0 - 0.3

Smaller number is the rate calculated with only impermeable-barrier boundaries and larger number is the rate calculated with a line-source boundary. See table 2.

#### North-Central Subarea

The north-central subarea includes a long, narrow section of the Anguilla Brook aguifer that extends about 1.7 miles north from Pequot Trail. It consists of relatively thin, saturated stratified drift (saturated thickness less than 20 feet) except near the southern end. Transmissivity ranges from 1,500 to over 3,000 ft $^2/d$  in the most favorable parts and the saturated thickness, in places, may be as much as 35 feet. The geologic contact between till and stratified drift, altitude of the bedrock surface, saturated thickness, transmissivity, assumed boundary conditions, and locations of hypothetical pumping wells are shown in figure 12. Six hypothetical wells with spacings that ranged from 800 to 1,750 feet were used to withdraw water for a 180-day period. Two different boundary configurations were assumed in evaluating this subarea; the first consists of two impermeable-barrier boundaries and one line-source boundary, while the second consists of three impermeable-barrier boundaries. The maximum long-term pumping rate calculated with the first set of boundary conditions was 0.7 Mgal/d, and the rate calculated for the second, more conservative set, was 0.6 Mgal/d. The amount of water estimated to be potentially available over the long term in this subarea is 1.8 Mgal/d if no upstream development occurs and 0.9 Mgal/d if it does. This means that the long-term ground-water yield in the north-central subarea, as indicated in tables 3 and 4, is between 0.6 and 0.7 Mgal/d, regardless of upstream development.



# MODEL BOUNDARIES AND HYPOTHETICAL WELLS

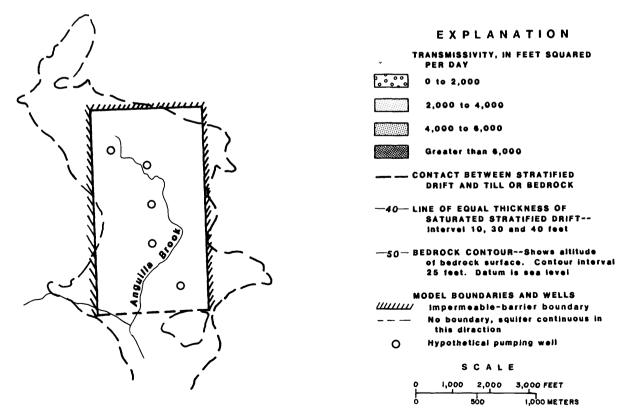
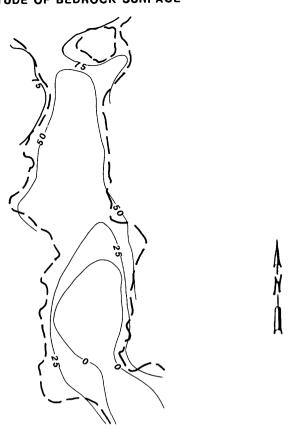
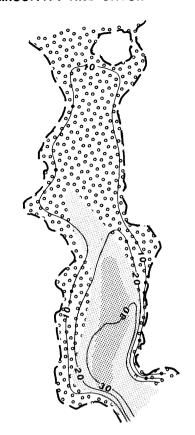


Figure 11.--Hydrogeologic features, assumed boundary conditions, and hypothetical pumping wells in the northern subarea of the Anguilla Brook aquifer.





#### MODEL BOUNDARIES AND HYPOTHETICAL WELLS

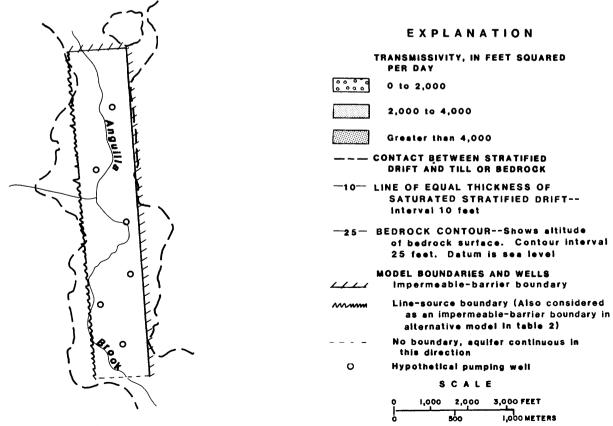


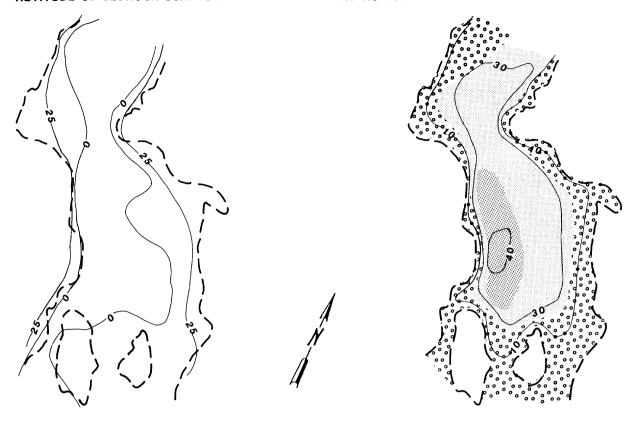
Figure 12.--Hydrogeologic features, assumed boundary conditions, and hypothetical pumping wells in the north-central subarea of the Anguilla Brook aquifer.

#### South-Central Subarea

The south-central subarea extends from Peguot Trail to U.S. Route 1. It contains coarse- to fine-grained stratified drift that tends to become finer with depth. Transmissivity ranges from 3.000 to about 3.700 ft<sup>2</sup>/d in the most favorable parts, and the saturated thickness, in places, may be as much as 40 feet. The geologic contact between till and stratified drift. altitude of the bedrock surface, saturated thickness, transmissivity, assumed boundary conditions, and locations of hypothetical pumping wells are shown in figure 13. Four hypothetical wells with approximately 1.000-foot spacings were used to withdraw water for a 180-day period and two boundary configurations were assumed. The first set of boundaries included two impermeable-barrier boundaries and one line-source boundary, while the second included three impermeable-barrier boundaries. The maximum, longterm pumping rate calculated with the first set of boundary conditions was 1.0 Mgal/d, and the rate calculated for the second set was 0.7 Mgal/d. The amount of water estimated to be potentially available over the long term in this subarea is 2.5 Mgal/d if no upstream development occurs and between 0.9 and 1.0 Mgal/d if it does. This means that the long-term ground-water yield in the south-central subarea, as indicated in tables 3 and 4, is 0.7 to 1.0 Mgal/d regardless of upstream development.

#### Southern Subarea

The southern subarea includes a small part of the Anguilla Brook aquifer, south of U.S. Route 1. Relatively little data is available from this subarea. Transmissivity is estimated to average about 2,500 ft<sup>2</sup>/d and saturated thickness may exceed 25 feet. The geologic contact between till and stratified drift, altitude of the bedrock surface, saturated thickness, transmissivity, assumed boundary conditions, and locations of hypothetical pumping wells are shown in figure 14. Two hypothetical wells, 750 feet apart, were used to withdraw water for a 180-day period. Two boundary configurations were assumed; the first consists of two impermeable-barrier boundaries and one line-source boundary, while the second has two impermeable-barrier boundaries. The maximum, long-term pumping rate calculated for both conditions was about 0.3 Mgal/d. The amount of water estimated to be potentially available over the long term is 2.6 Mgal/d if no upstream development occurs, and 0 to 0.4 Mgal/d if it does. The cumulative effect of upstream development is significant, and the long-term groundwater yield of the southern subarea, as indicated in tables 3 and 4, is 0.3 Mgal/d with no upstream development, and 0.3 Mgal/d or less with upstream development.



# MODEL BOUNDARIES AND HYPOTHETICAL WELLS

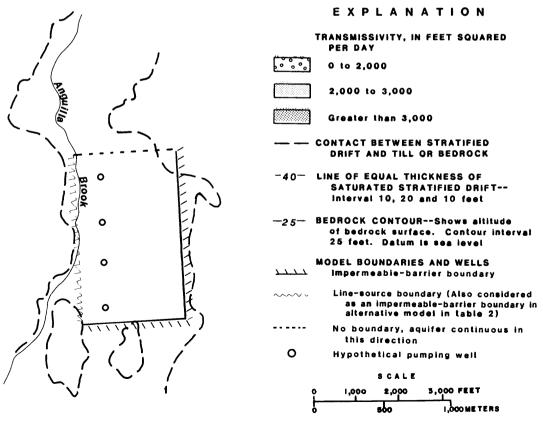
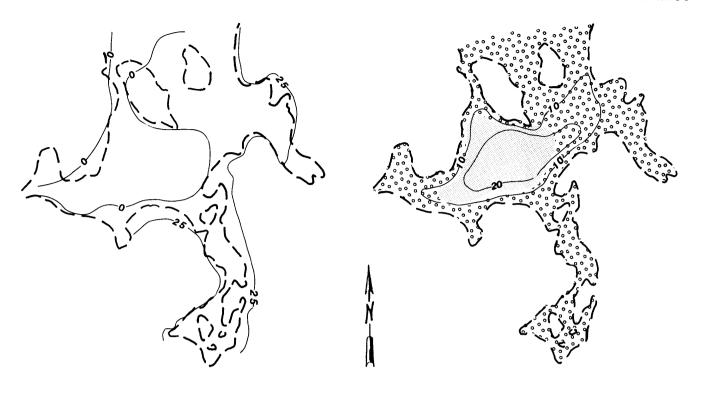


Figure 13.--Hydrogeologic features, assumed boundary conditions, and hypothetical pumping wells in the south-central subarea of the Anguilla Brook aquifer.



#### MODEL BOUNDARIES AND HYPOTHETICAL WELLS

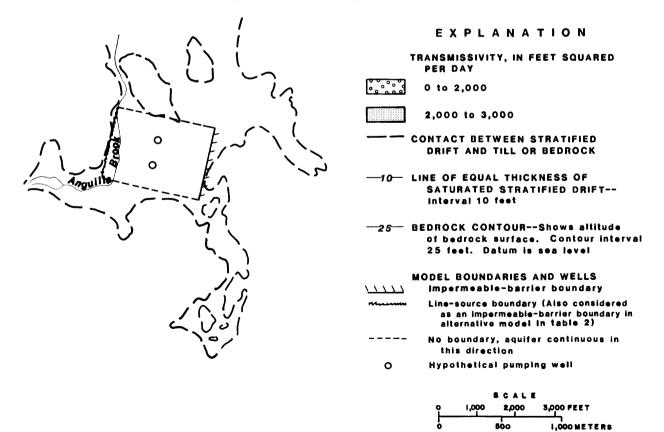


Figure 14.--Hydrogeologic features, assumed boundary conditions, and hypothetical pumping wells in the southern subarea of the Anguilla Brook aguifer.

# WATER QUALITY

## Surface Water

The water in Anguilla Brook and its tributaries was sampled and analyzed to see if it met drinking-water standards established by the State of Connecticut (Connecticut General Assembly, 1975), and the U.S. Environmental Protection Agency (1975). Table 5 includes the principal constituents that comprise the U.S. Environmental Protection Agency standards as well as the limiting values for drinking water. When a constituent exceeds the limiting value, the water is considered unsuitable for domestic supply without treatment.

Anguilla Brook was sampled at two sites, during average and low-flow conditions, to see if differences in water quality occur along the stream and as flow conditions vary. The sites were at Nathan Wheeler farm (station 01118535) where Anguilla Brook leaves the northern part of the basin and flows over bedrock outcrop, and at Wequetequock (station 01118550) near the downstream end of the basin. The concentrations of most constituents were greater at the downstream sampling site. Iron and manganese had the largest increase probably due to drainage from large areas of swampland in the central part of the basin.

Water samples were collected at both sites when discharges were representative of both average and low-flow (90-percent duration flow) conditions. Concentrations of most constituents were similar at average and low flows except for iron and manganese. Concentrations of these two dissolved metals decreased at low flows, probably due to reduced contributions of runoff from swamplands during low-flow periods. Water-quality data for the two sites, for samples collected at average (May 24, 1982) and low (August 19, 1982) flows are summarized in table 6. Analytical results for 55 organic compounds, including most common herbicides and pesticides, are not included in table 6 as none were present at the analytical detection limits. The organic compounds analyzed for are listed in Water Resources Data for Connecticut, Water Year 1982 (U.S. Geological Survey, 1982).

As the data indicate, water from Anguilla Brook exceeded recommended maximum limits for drinking water for only one constituent—dissolved iron, measured in the sample collected at Wequetequock (station 01118550) during average flow conditions. The dissolved iron concentration for this sample was 340  $\mu$ g/L (micrograms per liter), slightly above 300  $\mu$ g/L which is recommended as the maximum limit for this constituent.

#### Table 5 .-- Water-quality criteria

[The Federal Water Pollution Control Act Amendment of 1972 (P.L. 92-500) stipulated that water-quality criteria were to be developed to assure the integrity of ground and surface waters of the United States. Criteria were set for various types of water use. surface waters of the United States. Lriteria were set for various types of water These criteria indicate limiting values of various parameters in water to provide adequate protection of water users, essential aquatic life, and consumers of such aquatic life. Abbreviations: mg/L = milligrams per liter;  $\mu$ g/L = micrograms per liter; mL = milliliter; col/100 mL = colonies per 100 milliliters]

Parameter name	Limiting value	Units	<sub>Use</sub> <u>1</u> /	Basis for selection 2/	Parameter name	Limiting value	Units		Basis for selection
	General i	norganics				General inorgan	nics - cont	inued	
Alkalinity, total (as CaCO <sub>3</sub> )	<u>3</u> / 20	mg/L	2	Α	Selenium Silver	10 50	μg/L μg/L	4,6 4,6	A,B A,B
Arsenic	50 100	µg/L µg/L	4,6 3	A,B A	Solids, dissolved Sulfate	500 250	mg/l. mg,'L	6a 6a	C C
Barium Beryllium	1,000 11	μg/L μg/L	4,6 2a	A,B A	Zinc	5,000	µg/L	4,6a	A,C
_	100 1,100	μg/L μg/L	3 2b	A A			janics .		
Boron Cadmium	750 0.4 1.2	µg/L µg/L µg/L	3 la lb	A A A	Aldrin-dieldrin Chlordane	0.003 0.004 0.01	mg/L µg/L µg/L	2 8 2	A A A
	4.0 5.0 10	μg/L μg/L μg/L	2a 8 4,6	A A A,B	DDT <u>5</u> / Demeton Endosulfan	0.001 0.1 0.001	μg/L μg/L μg/L	2,8 2,8 8	A A
Chloride	12 250	μg/L mg/L	2b 6a	A C	Endrin	0.003	μg/L μg/L	2 2,8	Â A
Chromium, total	50 100	μg/L μg/L	4,6	A,B A	Guthion	0.2 0.01	μg/L μg/L	4,6 2,8	B A
Color	15 75	color units	6a 4	C A	Heptachlor Lindane	0.001 0.004	μg/L μg/L	2,8	A A
Copper Cyanide	1,000	μg/L μg/L	4,6a 2,8	A,C A		0.01 4	μg/L μg/L	2 4,6	A A,B
Fecal coliform, MF Iron	4/ 200 300 1,000	col/100 mL μg/L μg/L	7 4,6a 2	A A,C A	Malathion MBAS (foaming agents) Methoxychlor	0.1 0.5 0.03	µg/L mg/L µg/L	2,8 6a 2,8	A C A
Lead, dissolved Manganese Mercury	50 50 0.05 0.1	μg/L μg/L μg/L	4,6 4,6a 2 8	A,B A,C A A	Mirex Parathion	100 0.001 0.04	µg/L µg/L µg/L	4,6 2,8 2,8	A,B A A
Nickel Nitrate (as N)	2 100 10	μg/L μg/L μg/L mg/L	4,6 2,8 4,6	A A A A,B	PCB Phenols Toxaphene	0.001 1.0 0.005	µg/L µg/L µg/L	2,8 4 2,8	A A A
Nitrite (as N) Oxygen, dissolved pH	1 3/ 5 6.5-8.5 6.5-9.0	mg/L mg/L S	4,0 4 2 6a,8 2	A A A,C A	Silvex 2,4-D	5.0 10 100	μg/L μg/L μg/L	4,6 4,6 4,6	A,B A,B A,B
	5.0-9.0		4	Â					

- 1/ Water use and/or for the protection of:
  - la. Sensitive salmonoid species in soft water
  - 1b. Sensitive salmonoid species in hard water
  - Freshwater aquatic life
  - 2a. Freshwater aquatic life in soft water
  - 2b. Freshwater aquatic life in hard water3. Crop irrigation4. Domestic water supply source

  - 6. Potable drinking water, based on health effects
  - 6a. Potable drinking water, based on aesthetic considerations

  - Primary contact
     Marine aquatic life
- 2/ Basis for selection:
  - A. Maximum levels recommended by: Quality Criteria for Water, 1976, U.S. Environmental Protection Agency
  - B. Maximum contaminant level established by: National Interim Primary Drinking
  - Water Regulations, U.S. Environmental Protection Agency, 1975
    C. Maximum contaminant level recommended for the Proposed Secondary Drinking Water Regulations, U.S. Environmental Protection Agency, 1977

- 3/ Minimum recommended value
- 4/ Log mean, based on not less than five samples
- 5/ Including metabolites (DOD and DDE)

Table 6.--Analysis of surface-water samples from Anguilla Brook [<, less than; cfs, cubic feet per second;  $\mu$ S/cm, microsiemens per centimeter; °C, degrees Celsius; mg/L, milligrams per liter;  $\mu$ m, micron;  $\mu$ g/L, micrograms per liter; dashes indicate data not available]

	Station 011		Station 011	
	Anguilla Br		Anguilla Br	
	North Stoni		Wequetequoc	: <u>k</u>
Constituent or property	E (0E (00	Date of sample		0 /10 /02
	5/25/82	8/19/82	5/24/82	8/19/82
Streamflow, instantaneous (cfs)	8.6	1.4	11	2.5
Specific conductance (µS/cm)	81	89	87	98
oH (units)	6.7	6.3	6.8	6.5
「emperature (°C)	15.0	18.0		18.5
Color (platinum-cobalt units)	50	20	60	57
「urbidity (Nephelometric turbidity units)		1.0		2.0
Dxygen, dissolved (mg/L)	12.6	8.0		8.9
lxygen, dissolved (percent saturation)	123	84		95
Coliform, fecal, $0.45-\mu\mathrm{m}$ filter (colonies per 100 mL)		200		100
streptococci, fecal, KF agar (colonies per 100 mL)		190		320
lardness (mg/L as CaCO <sub>3</sub> )	22	25	24	28
Hardness, non-carbonate (mg/L as CaCO <sub>3</sub> )	7.0	10	5.0	10
Calcium, dissolved (mg/L as Ca)	5.7	6.5	6.4	7.5
Magnesium, dissolved (mg/L as Mg)	1.9	2.0	1.9	2.2
Sodium, dissolved (mg/L as Na)	6.0	5.9	6.3	7.3
Percent sodium	36	34	36	35
Sodium adsorption ratio	0.6	0.6	0.6	0.6
Potassium, dissolved (mg/L as K)	.7	.7	.8	.9
lkalinity (lab) (mg/L as CaCO <sub>3</sub> )	15	15	19	18
Sulfate, dissolved (mg/L as SO <sub>4</sub> )	10	9.0	8.0	9.0
Chloride, dissolved (mg/L as Cl)	9.4	9.2	10	10
Fluoride, dissolved (mg/L as F)	< .1	< .1	< .1	< .1
Silica, dissolved (mg/L as SiO <sub>2</sub> )	7.5	9.5	8.0	8.2
Solids, residue at 180 °C, dissolved (mg/L)	75	71	82	85
Solids, sum of constituents, dissolved (mg/L)	51	52	53	56
Nitrogen, NO <sub>2</sub> + NO <sub>3</sub> , dissolved (mg/L as N)	.38	. 57		.21
Phosphorus, dissolved (mg/L as P)	.010	< .010		< .010
luminum, dissolved ( $\mu$ g/L as Al)	50	30	90	40
Arsenic, dissolved (μg/L as As)	1	1	1	1
Barium, dissolved (µg/L as Ba)	42	39	32	32
Boron, dissolved (μg/L as B)	40	40	40	50
Cadmium, dissolved (µg/L as Cd)	1	< 1	< 1	< 1
Chromium, dissolved (µg/L)	< 1	< 1	< 1	< 1
Cobalt, dissolved (µg/L as Co)	4	< 1	3	< 1
Copper, dissolved (µg/L as Cu)	2	2	2	2
ron, dissolved (μg/L as Fe)	260	120	340	240
ead, dissolved (μg/L as Pb)	2	1	2	1
ithium, dissolved (μg/L as Li)	5	< 4	6	< 4
danganese, dissolved (μg/L as Mn)	22	7	37	20
dercury, dissolved (μg/L as Hg)	.1	< .1	.1	< .1
Nickel, dissolved (μg/L as Ni)	2	2	< 1	< 1
Selenium, dissolved (µg/L as Se)	< 1	< 1	< 1	< 1
Silver, dissolved (μg/L as Ag)	< 1	< 1	< 1	< 1
Zinc, dissolved (µg/L as Zn)	6	< 4	4	5
Cyanide, dissolved (mg/L as Cn)	< .01		< .01	
Methylene blue active substance (mg/L)	.03	.03	.02	.04

### Ground Water

Ground-water samples were collected from seven wells in the Anguilla Brook basin. Samples for analysis were obtained from each well at two different times (June and August, 1982) to insure that they accurately represented the quality of water in the aquifer. The results of the analyses of these water samples are given in table 7, and the locations of the sampled wells are shown on plate 1. Analytical results for 55 organic compounds are not included in table 7, as none were present at the analytical detection limits. The organic compounds analyzed for in the ground-water samples are also listed in the previously cited reference (U.S. Geological Survey, 1982).

Ground-water quality in the Anguilla Brook aquifer based on these samples indicates no evidence of contamination due to man's activities. Concentrations of dissolved manganese in five of the wells sampled (as much as 1,800  $\mu g/L$  in SN 172), and dissolved iron in one of the wells sampled (3,600  $\mu g/L$  in SN 174) exceeded limits recommended by the U.S. Environmental Protection Agency (1976) for domestic supply and are probably due to natural conditions. Excessive concentrations of dissolved manganese are evident only in newly constructed wells; chemical analyses of water samples from older wells and low-flow samples from Anguilla Brook show dissolved manganese concentrations below recommended limits. Except for the high concentrations of dissolved manganese and iron, none of the 14 ground-water samples analyzed exceeded limits for drinking water established by the State of Connecticut (Connecticut General Assembly, 1975) and established or recommended by the U.S. Environmental Protection Agency (1975, 1976).

## SUMMARY AND CONCLUSIONS

Estimates of the long-term ground-water yields from four areas of the Anguilla Brook aquifer are based on evaluations of the amounts of water potentially available and calculations of maximum pumping rates. The 90-percent duration flow of Anguilla Brook  $[0.4~({\rm ft^3/s})/{\rm mi^2}]$  is used as a measure of the amount of water potentially available from the aquifer. This flow value is dependent upon upstream drainage area and ranges from 0.9 Mgal/d for the northern (upstream) subarea to 2.6 Mgal/d for the southern (downstream) subarea.

Maximum long-term pumping rates are calculated with a mathematical model that considers boundary conditions, and well and aquifer characteristics. How effective a hydraulic boundary Anguilla Brook would be under conditions of development cannot be determined. Therefore, two boundary configurations are analyzed in the north-central, south-central, and southern subareas. One configuration assumes Anguilla Brook constitutes a line-source boundary; the other assumes only impermeable-barrier boundaries. This results in two values expressed as a range for long-term yield for these three subareas. The larger yield values represent the results of the analysis where Anguilla Brook is assumed to be a line-source boundary.

Two conditions of ground-water development are considered in calculating the long-term ground-water yields. The first condition treats each of the four subareas independently, and assumes that there are no upstream withdrawals. The long-term yields under this condition are 0.9 Mgal/d in the northern subarea, 0.6 to 0.7 Mgal/d in the north-central subarea, 0.7 to 1.0 Mgal/d in the south-central subarea, and 0.3 Mgal/d in the southern subarea.

The second condition assumes that development will take place throughout the aquifer and that any water withdrawn from an upstream subarea will be unavailable for downstream use. This reduces the water potentially available in three of the four subareas. The long-term yields under this condition are the same as in the first condition, with the exception of the southern area where there may be no water available because of the development of upstream areas. The total yield of all subareas is 2.6 Mgal/d and would require no reuse of water. Under both conditions, reductions in the flow of Anguilla Brook would occur, and in places, the stream channel could be dry as much as 10 percent of the time.

Analysis of water quality in the Anguilla Brook basin from both surface- and ground-water sources indicates that concentrations of all constituents, with the exception of dissolved manganese and iron, were below the drinking-water limits established by the State of Connecticut (Connecticut General Assembly, 1975) and established or recommended by the U.S. Environmental Protection Agency (1975, 1976).

Table 7. Analysis of ground-water samples from the Anguilla Brook aquifer [All wells tap stratified drift. Selected well characteristics listed in tables 8 and 9. Locations shown on plate 1. Symbols: <, less than; K, results based on colony count outside the accepted range. Abbreviations:  $\mu$ S/cm, microsiemens per centimeter;  $\mu$ m, microns; mL, milliliters; mg/L, milligrams per liter; °C, degrees Celsius;  $\mu$ g/L, micrograms per liter; dash indicates no data]

				ite ident	ification	n number
		N 69		164		165
	4125110	71523901	4123110	71522301	41224007	/1520801
	Date of	sample o	ollectio	n	<del></del>	
Constituent on property	6/2/02	0/10/02	6/1/02	0/17/02	612102	0 /17 /02
Constituent or property	6/3/82	0/10/02	0/1/02	8/17/82	6/2/82	8/17/82
Specific conductance (µs/cm)	96	102	-	102	134	130
pH (units)	-	5.7	-	5.7	-	5.7
Temperature (°C)	10.0	10.0	-	14.5	11.0	12.0
Color (platinum-cobalt units)	< 1		< 1	_	< 1	-
Coliform, fecal, 0.45 µm-filter (colonies per 100 mL)	< 1	< 1	-	< 1	< 1	< 1
Streptococci, fecal, KF agar (colonies per 100 mL)	< 1	2	-	84	50	4
Hardness (mg/L as CaCO <sub>3</sub> )	21	23	27	36	38	42
Hardness, non carbonate (mg/L as CaCO <sub>3</sub> )	10	15	13	23	24	31
Calcium, dissolved (mg/L as Ca)	5.9	6.6	7.0	9.2	11	12
Magnesium, dissolved (mg/L as Mg)	1.5	1.6	2.2	3.1	2.6	2.8
Sodium, dissolved (mg/L as Na)	5.6	7.1	2.3	2.9	4.8	5.6
Percent sodium	35	38	16	15	21	22
Sodium adsorption ratio	0.6	0.7	0.2	0.2	0.4	0.4
Potassium, dissolved (mg/L as K)	1.3	1.3	•5	•5	•9	.8
Alkalinity lab. (mg/L as CaCO3)	11	8.0	14	13	14	11
Sulfate, dissolved (mg/L as SO <sub>4</sub> )	12	12	10	9.0	19	18
Chloride, dissolved (mg/L as Cl)	9.0	9.7	4.4	9.0	12	11
Fluoride, dissolved (mg/L as F)	< .1	< .1	< .1	< .1	< .1	< .1
Silica, dissolved (mg/L as SiO <sub>2</sub> )	8.5	9.5	7.2	7.9	8.9	9.7
Solids, residue at 180°C, dissolved (mg/L)	62	80	68	103	84	98
Solids, sum of constituents, dissolved (mg/L)	51	53	42	50	68	67
Nitrogen, NO <sub>2</sub> + NO <sub>3</sub> , dissolved (mg/L as N)	1.4	1.3	2.5	2.8	.48	1.7
Phophorus, dissolved (mg/L as P)	< .010	< .010		< .010	< .010	< .010
Aluminum, dissolved (µg/L as Al)	20	10	20	20	30	30
Arsenic, dissolveds (µg/L as As)	< 1	1	< 1	1	< 1	1
Barium, dissolved (ug/L as Ba)	<b>4</b> 8	54	24	31	40	40
Boron, dissolved (ug/L as B)	<10	30	<10	30	30	50
Cadmium, dissolved (ug/L as Cd)	< 1	< 1	< 1	< 1	< 1	< 1
Chromium, dissolved (ug/L as Cr)	< 1	< 1	< 1	< 1	< 1	< 1
Cobalt, dissolved (ug/L as Co)	6	< 1	2	< 1	3	1
Copper, dissolved (ug/L as Cu)	< 1	59	< 1	22	1	2
Iron, dissolved (ug/L as Fe)	35	< 3	5	< 3	7	20
Lead, dissolved (ug/L as Pb)	2	< 1	< 1	2	1	2
Lithium, dissolved (ug/L as Li)	< 4	< 4	< 4	< 4	< 4	< 4
Manganese, dissolved (ug/L as Mn)	110	14	3	7	4	9
Mercury, dissolved (ug/L as Hg)	< .1	.1	< .1	.2	< .1	< .1
Nickel, dissolved (µg/L as Ni)	1	11	< 1	10	2	2
Selenium, dissolved (ug/L as Se)	< 1	< 1	< 1	< 1	< 1	< 1
Silver, dissolved (ug/L as Ag)	< 1	< 1	< 1	< 1	< 1	< 1
Zinc, dissolved (µg/L as Zn)	< 4	< 4	< 4	7	< 4	5
Cyanide, dissolved (mg/L as Cn)	< .01	_	< .01	_	-	-
Methylene blue active substance (mg/L)	.02	-	.03	-	-	-

Table 7. Analysis of ground-water samples from the Anguilla Brook aquifer--Continued

SN17 412443071		4124330715	SN172 412433071523901		SN174 412241071515401		22001
			8/18/82	6/2/82	8/17/82	6/3/82	8/17/82
84 6.8 10.0 <1 <1 <1 19 3.0 5.0 1.6 4.9 34 0.5 1.0 16 10 8.4 .1 6.5 50 48 .22 <.010 20 <1 <1 <1 <1 <1 <1 <1 20 4.9	8/18/82  87 6.5 16.0 - <1 1 20 2.0 5.2 1.7 6.9 41 0.7 1.2 18 9.0 7.7 .1 9.0 58 52 .12 <.010 30 1 29 40 <1 <1 <1 <1 2 24 2 <4 470 2	6/1/82  133 6.8 10.0 < 1 < 1 K1 39 0.00 11 2.8 6.1 24 0.5 1.7 41 13 7.7 2 9.3 80 78 -27 < .010 40 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10 < 1 34 10   1 34 10 10 10 10 10 10 10 10 10 10 10 10 10	8/18/82  120 6.4 11.5 - K3 2 36 1.0 10 2.7 6.2 26 0.5 1.4 35 11 7.5 .3 11 90 73 .49 <.010 10 1 33 30 <1 <1 <1 5 140 <1 <1 7,700 <1 2	6/2/82  185 6.9 11.0 < 1 < 1 57 5.0 15 4.7 8.1 23 0.5 2.8 52 17 10 .1 123 113 .48 < .010 90 < 1 130 10 1 < 1 5 < 1 2,700 < 1 440 .1 1	8/17/82  180 6.4 11.5 - <1 59 9.0 15 5.3 9.7 25 0.6 2.7 50 15 11 .1 25 127 118 .53 <.010 200 1 130 30 <1 1 2 2 3,600  1 <4 570 <1 1	6/3/82  128  11.0  < 1  < 1  < 1  < 1  6.0  25  0.5  2.0  27  15  8.0  < 1  15  84  75  1.8  < .010  20  < 1  23  10  < 1  < 1  < 4  450  < .1  2	8/17/82  125 6.3 12.5 - < 1 < 1 39 10 9.8 3.5 7.5 28 0.6 2.0 29 12 8.1 < 1 18 96 79 2.0 < .010 10 1 22 40 < 1 1 2 20 24 460 .2 3
< 1 < 1 < 4 < •01	< 1 < 1 5	< 1 < 1 6 < .01	< 1 < 1 < 4	< 1 < 1 9 < .01	< 1 < 1 11	< 1 < 1 < 4 < .01	< 1 < 1 6

Table 8.--Well and test-hole characteristics.

[Locations shown on plate 1. Well, test-hole, and site-location numbering systems explained in table 9.]

ber Site location number number and site of si		£	യയ	[ ] S	φ.	4	Date of		
20 4125150715250 01	Owner	Date drilled	NGVD of 1929)	land surface)	surface) $1/$	land surface)	water level measurement	il it Aquifer	Remarks <u>2</u> /
00 4195150715950 01				W E L	S				
20 4125160715250 01				North Stonington	nington				
69 4125110715239.01	E.J. Murphy & Sons Mrs. Jessie Lee	1958	105		35(R)	20.5	11-18-58 8-18-82	Stratified drift Stratified drift	Published in CWRB 16 L.O.Q.
NSN 70 4129170715229.01 NSN 71 4125130715211.01 NSN 73 4125210715322.01	koger Beaucage W.A. Clark, Est. Varian York	1977 1968 1976	115 140 160	27.5 52 405	25 22	35 35	4-6-77 8-29-68 9-29-76	Bedrock Bedrock Bedrock	L. 0-18 feet,till and boulders 0-22 feet,gravel and till
				Stonington	gton				
56 4121170715139.01	W.N. Siegel S. Zaromba Jr.	1 1	78	14.1		6.6	4-28-60	Lit	Published in CWRB 16.
4121420715124.01	Town of Stonington	•	52	95	•	. r. c	4-29-60	Bedrock	00.
75 4122220715111.01	ubaid kaveneile E. Morgan	1943	47	27	19	. 7 -	4-29-60	Bedrock	00°.
76 4122160715123.01	W.E. Medinger Town of Stonington	1955	£ %	9.5	15?	1.2	4-29-60	Stratified drift	Do.
80 4121560715157.01	Joseph Cangelosi	1959	888	180	22	10	1959	Bedrock	00.00
135 4121560715105.01	Town of Stonington	1962	38	10.7		0.8	8-20-63	Stratified drift	Do., backfilled from 13 feet
136 4121420/15124.02 141 4121520715205.01	lown of Stonington Thomas Bolton		22 53	32		ഹന	1959	Bedrock Till	Uo. Do.
142 4121560715157.02 143 4121430715225.01	Joseph Cangelosi Frank and Stanley Prachniak		20 15	18	19	3.7	1959 8-14-63	Stratified drift Stratified drift	Do.
145 4122190715113.01	Henry Richards		20	16.2		14.3	8-14-63		Do.
164 4123110715223.01 165 4122400715208.01	Town of Stonington Thomas F. Canaan Jr.	1976	84 %	19.6	44(R) 57(R)	7.34	8-17-82	Stratified drift	L.0.0.
170 4124520715228.01	Ralph Minor	1982	88	38.5	53(R)	3.62	6-2-82		.0.1
171 4124430715232.01 172 4124330715239.01	Alfred Minor Dudlev Wheeler	1982	80 74	30°0	42(R) 43(R)	3.57	6-1-82	Stratified drift	L.0.0.
173 4124330715223.01	Nathan Wheeler	1982	75	41.7	94 (R)	3.06	5-20-82		.0.1
174 4122410715154.01 175 4122520715220.01	Thomas F. Canaan, Jr. Carlton Cripps	1982 1982	41	30 <b>.</b> 0	32(R) 41(R)	3.82 6.25	5-18-82 5-17-82	Stratified drift Stratified drift	L.0.0.
179 4122510715222,01	Robert Cripps	1973	43	84	40	5	7-9-73		
180 4125070715223.01 181 4125070715331.01	Andrew Latham, Jr.	1972	235	30	, 30 27	12	2-18-72 8-19-68	Stratified drift Bedrock	2-inch diameter driven well.
183 4121270715151.01	noward chapman Town of Stonington	1965 1975	125 50	400	20	308	12-22-65 6-6 15	Bedrock Bedrock	Water supply for landfill.
			т Т	S	л 0 н	E S			
				Stonington	gton				
48 th 4123330715229.01 49 th 4123310715225.01	Conn. DOT	1962 1962	47	16.5	11.5	00	1962 1962	Stratified drift Stratified drift	L. Deepest of 4 holes. L. Deepest of 4 holes.
136 th 4122250715157.01	Thomas Canaan, Jr.	1976	30	50	>50	· 20	9-30-76	Stratified drift Stratified drift	-
139 th 4124230715245.01	Dudley Wheeler	1982	24.5	24	24(R)	۱ m r	4-21-82	Stratified drift	· .
th 4122270715152,01	Nathan wheeler Thomas Canaan, Jr.	1982	32	44.0	44.5(K) 43(R)	<b>4</b>	4-23-82	Stratified drift Stratified drift	
142 th 4124060715232.01	Edward Davis	1982	89	25	25(R)	10	4-30-82	Stratified drift	۲.

<sup>1/</sup> Depth to bedrock: (R), Refusal. Penetration of auger flight stopped by boulder or bedrock.

2/ Remarks:

CWRB 16, Connecticut Water Resources Bulletin 16 (Cervione and others, 1968) L., Log in table 9, 0. Observation well, Q., Quality of water analysis in table 7.

#### Table 9.--Logs of wells and test holes.

[Under each town, the logs are listed by their town well or test-hole number followed by the location number (latitude and longitude), owner, altitude, the year drilled where available, driller, depth to water, length of screen and depth that bottom of screen is set, and measuring point for observation wells in feet above land surface datum (LSD)]

Well and test-hole identification and site-location numbers: U.S. Geological Survey number assigned to each site. The letter prefix denotes the town in which it is located followed by a sequential number. The test holes are identified by the "th" suffix. Location number is the latitude and longitude. Number after decimal point is a sequential number used to identify closely spaced wells and test holes.

Altitude: The land surface at the site, in feet above NGVO of 1929, estimated from a topographic map with a contour interval of 10 feet, except at Connecticut Department of Transportation sites where land surface altitude is determined by leveling.

Depth to water: Expressed in feet below land surface.

Description of earth materials: The descriptive terms are those of the driller or geologist; logs of test holes of the U.S. Geological Survey and of the Connecticut Department of Transportation are based on the corresponding grain-size classification shown in the table at the right. Some Connecticut Department of Transportation logs use terms that approximate the percentage of a grain size within the described interval as follows:

	rercer
trace :	0-10
little:	10-20
some :	20-35
and:	35-50

Grain ize in illimeter (mm)		grade scale for ogical Survey logs	Grade scale used by Conn. Dept. of Trans- portation before 1959	AASHO Classification used by Conn. Dept. Transportation since about 1959	of .
255		Boulders		Boulders	
64		Cobbles		Cobbles	-203 mm
32		Very coarse gravel	1		— 76.2 mm —
16	Pebbles	Coarse gravel	Gravel	Medium gravel	— 25.4 mm —
в	_	Medium gravel		Fine gravel	
4		Fine gravel		, ,,,,	
2	Granules	Very fine gravel	3		2 mm
-	Very coarse	sand	2 mm ——		— 2 mm —
1	Coarse sand	i	Coarse sand	Coarse sand	
0.5	Medium sand	1	Medium sand		— 0.42 mm—
0.25	Fine sand		0.2 mm	Fine sand	
0.125	Very fine s	sand	Fine sand	5414	0.074 mm-
0.062	Silt		0.06 mm	Silt	0.004
0.004	Clay		Clay 0.002 mm	Clay	0.004 mm -

NSN 69. 4125110715239.01. Mrs. Jessie Lee. Altitude 112 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 20.5 ft on 8-18-82. Set 4.8 ft of slotted screen at 35.0 ft. Measuring point: Top of PVC casing 1.75 ft ab

Material Description	Depth (ft below LSD) From To	Thick- ness (ft)
Soil, sandy loam	0 - 1	1
Gravel, granule to boulder, and sand, very coarse; little sand, very fine to medium; trace of silt		8
to cobble, little silt and very fine sand; fewer cobbles with depth		26

NSN 70. 4125170715229.01. Roger Beaucage. Altitude 115 ft. Drilled 1977 by Carl Anderson. Depth to water 25 ft on 4-6-77.

	(ft bel	th ow LSD)	ness
Material Description	From	То	(ft)
Gravel, dry	12	- 25	12 13 209
Seam, water-bearing	234	-235	1
Ledge, soft	235	-275	40

SN 164. 4123110715223.01. Town of Stonington. Altitude 48 ft. Drilled 1976 by U.S. Geological Survey. Depth to water 7.34 ft on 8-17-82. Set 5 ft of screen at 19.6 ft. Measuring point: Top of PVC casing 0.3 ft above LSD.

Material Description	Dep (ft be From			Thick- ness (ft)	
					-
Fill	. 0	-	2	2	
Gravel, medium, and sand, fine to coarse, littl	e				
silt	. 2	- 1	10	8	
Sand, medium to coarse	. 10	- 1	13	3	
Sand, coarse to very coarse, little gravel,					
medium	. 13	- 1	17	4	
Sand, coarse to very coarse, and gravel, fine .				10	
Sand, medium to coarse, some very coarse				5	
Sand, very fine to fine, some silt		- 3	39	7	
Till, sandy, gravelly, and silty (dirty sand an					
gravel)				5	
Refusal	. a1	t 44	ļ.		

SN 165. 4122400715208.01. Thomas F. Canaan, Jr. Altitude 36 ft. Drilled 1976 by U.S. Geological Survey. Depth to water 2.79 ft on 5-18-82. Set 5 ft of screen at 22.2 ft. Measuring point: top of PVC casing 1.7 ft above LSD.

	De	pti	1	Thick-
	(ft be			ness
Material Description	From		10	(ft)
Topsoil	0	-	2	2
little silt	. 2	_	7	5
Sand, medium to coarse, clean	. 7	-	17	10
Sand, fine to medium, some very fine sand	. 17	-	30	13
Sand, very fine, some fine sand	30	-	37	7
Sand, very fine, and silt, some clay	37	-	49	12
pebbles		t !		8

. SN 170. 4124520715228.01. Ralph Minor. Altitude 80 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 3.62 ft on 6-2-82. Set 2 ft of screen at 38.5 ft. Measuring point: top of PVC casing, 2.9 ft above LSD.

Material Description	(ft below L	
Silt, light brown		
Sand, fine to very coarse, and pebble gravel. Sand, fine to very coarse, layered, brown;	7 - 20	13
occasional layers of silt or granule gravel . Sand, very fine to very coarse, and granule to	20 - 38	3 18
pebble gravel, trace of silt, firm, layered . Refusal		3 15

SN 171. 4124430715232.01. Alfred Minor. Altitude 80 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 3.57 ft on 6-1-82. Set 4.8 ft of screen at 39.0 ft. Measuring point: top of PVC casing 2.8 ft above LSD.

Material Description	Depth (ft below LSD) From To	ness (ft)
Soil, silty loam, black		2
silt, trace of granule gravel	. 2 - 21	19
iron-stained	. 21 - 25	4
trace of very fine sand		13
Sand, medium to very coarse, and pebble gravel	. 38 - 42	4

SN 172. 4124330715239.01. Dudley Wheeler. Altitude 74 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 2.49 ft on 5-20-82. Set 4.8 ft of screen at 30.4 ft. Measuring point: top of PVC casing 2.0 ft above LSD.

Material Description			h w LSD) To	Thick- ness (ft)
Soil, silty, and sandy	. 0	-	4	4
Sand, very fine to very coarse, silty, light brown	. 4	-	8	4
granule gravel, trace of very fine sand, rusty brown		-	18	10
very coarse	. 18	-	26	8
Gravel, granule to pebble, and sand, fine to very coarse, trace of silt		-	29	3
some fine to medium sand		-	39	10
Sand, medium to very coarse, granule to pebble gravel, trace of silt and very fine sand Refusal		at ·		4

SN 173. 4124330715223.01. Nathan Wheeler. Altitude 75 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 3.06 ft on 5-20-82. Set 4.8 ft of screen at 41.7 ft. Measuring point: top of PVC casing, 2.6 ft above LSD.

	D <sub>i</sub>	epti	1	Thick-
	(ft b	e lo	W LSD)	ness
Material Description	Fro	n	To	(ft)
Soil, gravelly sand	. 0	-	0.5	0.5
to very coarse, little silt	. 0	.5-	7	6.5
granule to pebble	. 14	-	14 27	7 13
occasional layers of well-sorted very fine san and silt, or medium to very coarse sand Sand, very fine to fine, laminated, clean:		-	42	15
layered with fine to medium sand			52 53	10 1
Sand, medium to very coarse	. 53	-	56	3
Sand, fine to medium, with silt layers Gravel and sand			64 68	8 <b>4</b>
Sand, fine to coarse			90 94	22 4
Refusal		at 9	94	

SN 174. 4122410715154.01. Thomas F. Canaan, Jr. Altitude 38 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 3.82 ft on 5-18-82. Set 4.5 ft of screen at 30.0 ft. Measuring point: top of PVC casing 2.0 ft above LSD.

Material Description	(ft below LSD) From To	ness (ft)
Soil, silt, black	. 0 - 2	2
granule to small pebble, some silt		5
brown		7
Sand, very fine to fine, and silt, gray	14 - 25	11
granule to large pebble		7

SN 175. 4122520715220.01. Carlton Cripps. Altitude 41 ft. Drilled 1982 by U.S. Geological Survey. Depth to water, 6.25 ft on 5-17-82. Set 3.4 ft of screen at 20.8 ft. Measuring point: top of PVC casing, 2.0 ft above LSD.

Material Description	Depth (ft below LSD) From To	ness (ft)
Soil, silt, black	. 0 - 1	1
Gravel, granule to large pebble, sand, very fine to very coarse, and silt	1 - 12	11
to small pebble	12 - 21	9
Sand, very fine to fine, gray, very soft Silt, gray; laminated, interbedded with thin		10
layers of medium sand		4
fine to very coarse; trace of silt	35 - 41	6

#### TEST HOLES

SN 48 th. 4123330715229.01. Conn. Department of Transportation. Drilled 1962. Altitude 47 ft. Water level 0 ft. Three additional borings (not shown) 11.5 to 15 ft deep, at this site.

Depth Thick-

( Material Description	ft below LSD) From To	ness (ft)
Water	0 - 2	2
Sand, fine to coarse, gray, and fine to coarse gravel; little silt; trace of clay		3
Sand, coarse to fine, brown to gray, and fine to coarse gravel; little silt	. 5 - 11.5	6.5 2
Gneiss. dark-gray, micaceous	13.5- 16.5	2 3

SN 49 th. 4123310715225.01. Conn. Department of Transportation. Drille 1962. Altitude 46.5 ft. Water level 0 ft. Three additional borings (not shown), 11.5 to 15 ft deep, at this site.

Material Description	(ft below From	w LSD)	ness (ft)
Water		0.5	0.5
Sand, coarse to fine, brown, and fine to coarse gravel; trace of silt	. 0.5-	8.5	8
gravel; trace of silt		12.5	4
to coarse gravel; trace of silt	. 12.5-	18.5	6

SN 136 th. 4122250715157.01. Thomas Canaan, Jr. Altitude 30 ft. Drilled 1976 by U.S. Geological Survey. Depth to water, 5 ft.

Material Description	Depth (ft below LSD) From To	Thick- ness (ft)
Topsoil Sand, coarse to very coarse, some gravel Sand, coarse to very coarse Sand, fine to coarse, some silt, gray Silt, gray, some clay Sand, very fine to fine, some silt and clay Gravel, fine to coarse, and sand, medium to ver coarse, silty; some layers of very clean fine	2 - 7 7 - 10 10 - 13 13 - 25 25 - 36	2 5 3 3 12 11
sand	. 45 - 50	9 5

SN 138 th. 4124280715252.01. Dudley Wheeler. Altitude 76 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 2 ft.

Material Description		ept lo	h w LSD) To	Thick- ness (ft)
		_		
Soil, silty, black	. 0	-	2	2
Gravel, pebble to cobble, and sand, fine to				
very coarse, some silt, brown	. 2	-	4	2
Sand, fine to very coarse, some silt		-	7	3
Gravel, pebble, and sand, fine to very coarse .		-	16	9
***********		_	17	1
Gravel, granule to cobble, and sand, fine to very coarse; with layers of gray silt and very				
fine sand		_	27	10
Sand, fine to very coarse, some silt, gray, ver	y			
hard		-	32	5
Weathered bedrock (gneiss)	. 32	-	36	4
Bottom of hole		t	36	

SN 139 th. 4124230715245.01. Dudley Wheeler. Altitude 74 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 3 ft.

Material Description	De ft be From			Thick- ness (ft)
Soil, silty, black	. 0	-	2	2
Gravel, granule to small pebble, and sand, medium to very coarse, with some very fine sand between 5-7 ft	1			13
Sand, very fine to very coarse, silty; few pebbles, firm	. 18			5 1
Refusal		2		•

SN 140 th. 4124250715237.01. Nathan Wheeler. Altitude 79 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 7 ft.

	Depth	Thick-
Material Description	(ft below LSD) From To	ness (ft)
Soil, loam, black	0 - 1	1
very coarse, little very fine sand		6
to pebble, trace of fine sand		15
pebble		5
Sand		3
Sand, very fine to very coarse, silty, and		5
gravel, granule to pebble, firm		9

# Table 9.--Logs of wells and test holes--Continued

SN 141 th. 4122270715152.01. Thomas F. Canaan, Jr. Altitude 32 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 4 ft.

Material Description		oth low LSD) To	Thick- ness (ft)
Soil, sandy silt, brown	. 0	- 1	1
very coarse	. 1	- 10	9
Gravel, granule, and sand, fine to very coarse		- 13	3
Sand, very fine to very coarse	. 13	- 24	11
Silt, gray and brown, laminated	. 24	- 30	6
with clay layers; some cobbles	. 30	- 40	10
Till, sandy, silty, brown, very hard Refusal	. 40	- 43 : 43	3

SN 142 th. 4124060715232.01. Edward Davis. Altitude 68 ft. Drilled 1982 by U.S. Geological Survey. Depth to water 10 ft.

Material Description		th ow LSD) To	Thick- ness (ft)
Carral annuals to salely and a first			
Gravel, granule to cobble, sand, very fine to very coarse, and silt, gray-brown	0	20	20
Till, clayey, silty; yellowish brown, some	. 0	- 20	20
granules to pebbles; firm	20	_ 23	3
Bedrock, weathered			2
Refusal		25	_

Table 10.--Grain-size analyses of samples of stratified drift from the Anguilla Brook aquifer

[All samples are disturbed but uncontaminated. They were collected from a split-spoon sampler or auger flight during U.S. Geological Survey drilling. The well and test-hole locations are shown on plate 1 and the logs are in table 9. All analyses and calculations were made by the U.S. Geological Survey. Well, test-hole, and site-location numbering systems explained in table 9. Abbreviations: ft, feet; mm, millimeter; in., inch]

Depth sampled: Interval in feet below land surface from which sample was taken.

Grain-size distribution in percent of total weight finer than given size. Size intervals are those of the Wentworth scale (shown at the beginning of table 9).

Median grain size: The grain size at which 50 percent of the material is coarser and 50 percent finer. The size is read from a graph at the midpoint of the cumulative distribution curve.

Percent of total weight finer than given size													
Well or test-hole	· · ·		Median	32 mm	16 mm	8	<b>4</b> <b>m</b> m	2 	1 mm	0.5 mm	0.25 mm	0.125 mm	0.062 mm
identifi- cation number	Site location number	Depth sampled (ft)	grain size (mm)		Gr	avel		Very coarse sand	Coarse sand	Medium sand	Fine sand	Very fine sand	Silt and clay
- Maniber	Hamber												
					W	E L	L S						
				Town o	f North	Stoningt	on						
NSN 69	4125110715239.91		3.8	100	84	64	52	43	35	27	19	12	7
		31-33	0.6	100	96	88	80	73	61	46	32	22	13
				Tow	n of Sto	nington							
SN 164	4123110715223.01	17-18.5 22-23.5	0.8 1.8		100	100 91 <b>.</b> 1	95.8 67.5	82.1 51.7	63.8 41.6	25.6 26.5	10.7 9.4	3.8 3.1	1.9 1.1
		27-28.5	0.75		100	95.6	86.6	76.3	64.2	32.9	12.2	3.2	1.4
		32-33.5	0.14			100	99.3	98.7	97.0	93.3	77.7	40.6	19.5
		(top 3 in.) 32-33.5	0.06					100	99.9	99.6	98.3	84.4	53.3
		(bottom 3 in 42-43.5	.)			100	99.2	96.2	86	75.6	64.5	43.4	26.3
		(top 6 in.)							-				
		42-43.5 (bottom 6 in	0.3 .)		100	91.2	82.1	71.8	64.1	57.7	48.6	36.0	25.1
SN 165	4122400715208.01	22-23.5	0.22		100	99.5	95.6	94.0	93.0	88.7	53.4	19.2	7.2
		27-28.5 37-38.5	0.08 0.6					100 100	99.8 99.9	99.6 99.8	97.7 99.6	67.7 91.4	32.3 49.6
SN 170	4124520715228.01	17-18 22-24	1.8 0.3	100	91 100	76 99	63 94	53 90	43 79	33 62	24 42	16 21	10 11
		27 - 28	0.3		100	96	89	84	77	62	45	29	19
		32-34	0.6		100	95	84	73	61	46 36	31 27	18 18	10 11
		37-39 52-53	1.4 0.4		100 100	86 96	67 92	56 8 <b>6</b>	45 73	56	41	24	12
SN 171	4124430715232.01	10-15	0.6		100	98	93	88	72	44	18	8	4
		22-24	0.5	100	100	83	74 60	68	61 42	51 29	40 20	29 12	15 6
		27-29 32-34	1.8 0.6	100 100	82 96	70 91	82	52 74	42 61	29 46	20 35	20	9
		37-38	0.6	100	100	92	77	68	57	47	39	20	8
		38-39	4.5	100	92	63	46	39	33	28	24	16	7
SN 172	4124330715239.01	4- 6 9-12	0.8 0.7		100 100	91 88	79 77	68 69	56 57	40 41	17 17	8 7	4 4
		22-24	3.0		100	84	58	42	29	19	12	6	3
		27-29	1.1		100	83	64	56	49	42	23	15	3
		32-34 37-39	1.1 0.8	100	100 96	85 82	77 72	64 64	49 55	30 44	17 32	5 21	1 12
		42-43	0.3	100	100	96	87	79	71	61	49	34	22
SN 173	4124330715223.01	16-18	0.3					100	98	76	39	21	12
		21-23	0.2				100	100	99	91	61	23 22	8 13
		26-28 31-33	0.5 0.6			100	100 9 <b>5</b>	78 87	63 67	50 44	<b>36</b> 26	12	6
		36-38	0.5		100	99	87	79	68	49	29	15	8
		41-43	0.13						100 100	99 99	90 96	42 61	13 22
		46-48 51-53	0.10 0.05						100	100	99	96	65
		5 <b>6-5</b> 8	0.20			100	<b>9</b> 9	95	85	70	<b>5</b> 8	28	9
		61-63 76-79	0.07 0.11						100 100	99 96	94 88	76 53	44 17
		10-19	0.11						100	20	00	33	• /

Well or				32	16		ercent o	2	1	ner than	0.25	0.125	0.062
est-hole dentifi-	Site	Donth	Median	mm	ma		mm_	Min	mm	mm	mm	mm	mm
cation	location number	Depth sampled (ft)	grain size (mm)		Gr	avel		Very coarse sand	Coarse sand	Medium sand	Fine sand	Very fine sand	Silt and clay
					W	E L	L S						
					Town of	Stonin	gtonCo	ntinued					
SN 174	4122410715154.01	3- 6 8-11 18-20 21-23 31-32	0.8 2.6 0.08 0.03	100 100	92 88 100	79 71 80	71 57 65	64 46 100 58	54 33 99 50	40 18 98 100 41	28 10 91 99 31	21 7 69 97 20	16 6 39 81 12
SN 175	4122520715220.01	5- 6 10-11 16-18 26-31	1.5 2.9 0.6 0.06	100	100 76 100	83 61 89	66 53 79	55 47 71	45 40 61	36 31 48	25 23 30 100	15 16 14 90	9 12 8 42
		31-33 36-38	0.01 1.5	100	92	71	62	55	46	100 37	99 26	98 16	94 10
				T	E S	T	н 0	L E	S				
					<u>To</u>	wn of S1	oningto	<u>n</u>					
SN 138 t	h 4124280715252.0	11- 13 16 -17 17.3-18 21- 23 26- 27 31- 32	1.2 0.3 4.5 0.8 0.5	100 100 100	97 77 96 100	58 81	65 47 72 84 96	58 100 40 64 69 92	47 98 32 53 59 82	35 75 24 42 50 64	23 30 17 29 35 36	14 6 11 19 20 14	8 2 6 11 10 5
SN 139 t	h 4124230715245.0	11 16- 18 21- 23	1.1		100	79 100	67 93	59 88	49 81	39 67	28 46	18 26	11 15
SN 140 t	h 4124250715237.0	11 18- 20 22- 24 26- 28 31- 32 36- 38 42- 43	0.5 0.8 0.3 0.09 1.0 0.21	100	100 100 87 100	93 100 78	82 85 99 65 85	76 73 96 58 81	65 57 88 100 49 74	51 39 72 99 40 65	37 26 47 97 30 54	20 16 22 76 21 40	9 8 11 25 13 26
SN 141 t	h 4122270715152.0	11 4- 6 8- 10 14- 16 18- 20 23- 25 31- 33	0.7 3.5 3.5 0.10 0.02 1.5	100 100 100	93 78 90	65 65 100	69 52 53 98	63 46 45 95 100 52	56 39 36 90 99 45	44 31 26 80 98 38	28 20 14 75 97 29	14 11 8 61 95 20	8 7 5 24 76 14
		41- 42.		100	95 100	80	72 72	65 66	58 59	50 49	38 38	25 2 <b>6</b>	14 15
SN 142 t	h 4124060715232.0	1 8- 16 20- 23	1.8 0.08	100	98 100		60 92	52 88	43 85	32 78	22 66	14 54	8 42

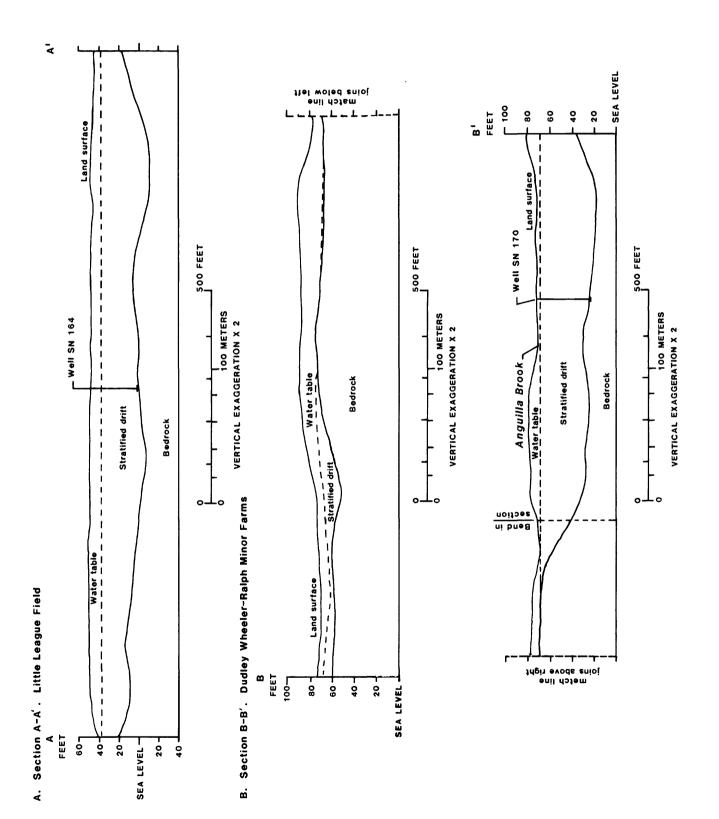
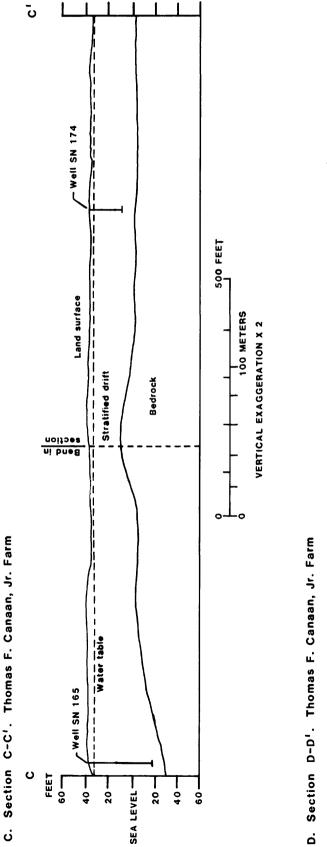
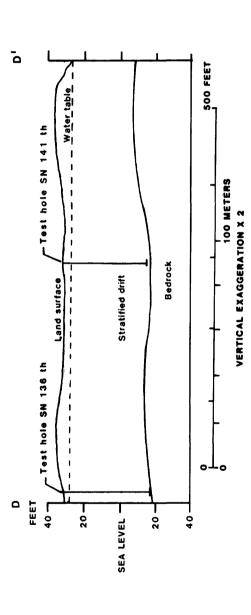
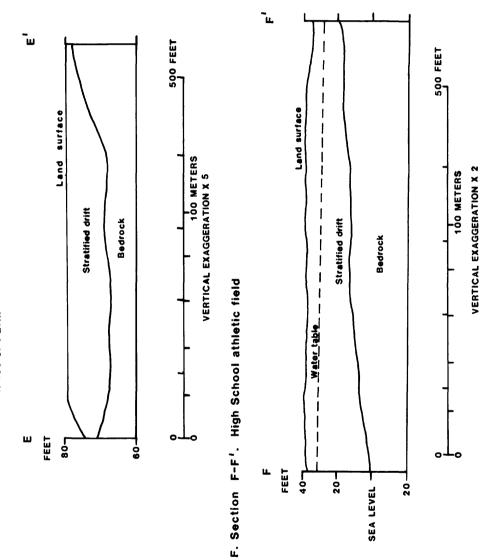


Figure 15.--Cross sections showing hydrogeologic units as interpreted from seismic-refraction data in Anguilla Brook valley, Stonington, Connecticut.





E. Section E-E'. Nathan Wheeler Farm



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